

**Research** Article

# Variation in Argument of Perigee for Near-Earth Satellite Orbits Perturbed by Earth's Oblateness and Atmospheric Drag Interms of Ks Elements

#### Lila S Nair\*

Department of Mathematics, HHMSPB NSS College for Women, Thiruvananthapuram, India

#### Abstract

Analytical solutions with the KS element equations of motion due to the combined effect of zonal harmonics J2,J3 and J4 and drag by considering an analytical oblate diurnal exponential density model when density scale height varies with altitude is obtained using series expansion method. Terms up to third terms in e, eccentricity, c, a small parameter depending on the ellipticity of the atmosphere and second order terms in  $\mu$ , gradient of the scale height altitude are considered. The KS element equations are numerically integrated (NUM) through a fixed step size fourth-order Runge-Kutta-Gill method having a very small step-size of half degree in the eccentric anomaly for comparing analytically integrated (ANAL) values. After 100 revolutions, decrease in argument of perigee,  $\omega$ , at perigee height = 400 kilometer, e = 0.1 and inclination i = 20 and 80 degrees, are found to be 7.42 and 39.8 degrees. At i =80 degree, the percentage error = (ANAL - NUM) / NUM after 1 and 100 revolutions are 0.61 and 2.09.

**Keywords:** Ks elements; Zonal harmonics; Atmospheric drag; Analytic integration

### Introduction

The dynamical system of a satellite motion perturbed by both atmospheric drag and gravitational attraction is nonlinear, non conservative in form and the integration of the system, in general, is analytically intractable. To predict the motion precisely a mathematical representation for these forces must be selected for integrating the resulting differential equations of motion. Some of the early studies and analytical difficulties for the coupled problem were addressed by de Nike [1]. Hoots [2] used the gravitational and atmospheric models as used by Lane [3] and arrived at an improved analytical solution. Well known and commonly used models [4-9]. The KS total-energy elements equations [10] is a very powerful method for numerical solution with respect to any type of perturbing forces as the equations are less sensitive to round-off and truncation errors in the numerical integration algorithm. Sharma worked with these KS element equations to compute very accurate short-periodic terms due to  $J_{2}$ , even for very high eccentricity orbits [11,12]. Sharma [13,14] expanded analytic solutions by series expansion method using analytical models for oblate exponential atmospheric density model, and a model of the same with the effect of diurnal bulge.

In this paper my attempt is to get analytical solutions with the KS element equations of motion for long-term motion by considering the perturbations due to the combined effect of Earth's zonal harmonics  $J_2$  to  $J_4$  and atmospheric drag. The model used here is an oblate diurnally varying atmosphere with variation of the scale height depending on altitude almost similar to my work with Sharma [15]. Using series expansion method third-order terms in e, eccentricity, c, a small parameter depending on the ellipticity of the atmosphere and second order terms in  $\mu$ , gradient of the scale height altitude are collected. Only one of the eight equations is solved analytically to obtain the state vector at the end of each revolution due to symmetry in KS equations. Numerical studies with test cases reveal that there is a good comparison between the analytical (ANAL) and numerically integrated (NUM) values of the position as well as velocity vectors  $\vec{x}$  and  $\vec{x}^*$ .

# **Equations of Motion**

The KS element equations of motion of a satellite under the effect

of perturbing potential V and additional perturbing force  $\vec{P}$  [10] are

$$\frac{d\omega}{dE} = -\frac{r}{8\omega^2} \frac{\partial V}{\partial t} - \frac{1}{2\omega} \left( \frac{d\vec{u}}{dE}, L^T(\vec{u})\vec{P} \right).$$
(1)

$$\frac{d\tau}{dE} = \frac{1}{8\omega^3} \Big[ K^2 - 2rV \Big] - \frac{r}{16\omega^3} \left( \vec{u}, \frac{\partial V}{\partial \vec{u}} - 2L^T(\vec{u})\vec{P} \right) - \frac{2}{\omega^2} \frac{d\omega}{dE} \left( \vec{u}, \frac{d\vec{u}}{dE} \right).$$
(2)

$$\frac{d\vec{\alpha}}{dE} = \left\{ \frac{1}{2\omega^2} \left[ \frac{V}{2} \vec{u} + \frac{r}{4} \left( \frac{\partial V}{\partial \vec{u}} - 2L^T(\vec{u})\vec{P} \right) \right] + \frac{2}{\omega} \frac{d\omega}{dE} \frac{d\vec{u}}{dE} \right\} \sin \frac{E}{2}.$$
(3)

$$\frac{d\vec{\beta}}{dE} = -\left\{\frac{1}{2\omega^2}\left[\frac{V}{2}\vec{u} + \frac{r}{4}\left(\frac{\partial V}{\partial \vec{u}} - 2L^T(\vec{u})\vec{P}\right)\right] + \frac{2}{\omega}\frac{d\omega}{dE}\frac{d\vec{u}}{dE}\right\}\cos\frac{E}{2}.$$
(4)

 $K^2 = k^2(M + m)$ , specifying attraction between two masses M and m, E,  $\omega, t, r$  and  $k^2$  are, respectively, the eccentric anomaly, angular frequency, physical time, radial distance and the gravitational constant.

The perturbing potential V[10] and the aerodynamic drag force  $\vec{P}$  [16] per unit mass acting on a satellite of mass m are respectively

$$V = \frac{K^2}{r} \sum_{n=2}^{\infty} J_n \left(\frac{R}{r}\right)^n P_n \cos(\nu) \text{ where } \cos(\nu) = \frac{x_3}{r},$$
(5)

R, equatorial radius,  $J_n$ 's, dimensionless constants known as zonal harmonics. Using Equation (5),

$$V_{2} = \frac{K^{2}J_{2}R^{2}}{2r^{3}} \left[ -1 + 3\frac{x_{3}^{2}}{r^{2}} \right]; \qquad V_{3} = \frac{K^{2}J_{3}R^{3}}{r^{4}} \left[ -\frac{3}{2}\frac{x_{3}}{r} + \frac{5}{2}\left(\frac{x_{3}}{r}\right)^{3} \right];$$
$$V_{4} = \frac{K^{2}J_{4}R^{4}}{r^{5}} \left[ \frac{3}{8} - \frac{15}{4}\left(\frac{x_{3}}{r}\right)^{2} + \frac{35}{8}\left(\frac{x_{3}}{r}\right)^{4} \right]; \quad \left(\vec{u}, \frac{\partial V}{\partial u}\right) = -2(n+1)V_{n} \text{ and } \frac{\partial V}{\partial t} = 0,$$
where

\*Corresponding author: Lila S Nair, Department of Mathematics, HHMSPB NSS College for Women, Thiruvananthapuram, India, Tel: 9847003427; E-mail: lilasnair@yahoo.co.in

Received October 02, 2015; Accepted October 15, 2015; Published October 25, 2015

Citation: Nair LS (2015) Variation in Argument of Perigee for Near-Earth Satellite Orbits Perturbed by Earth's Oblateness and Atmospheric Drag Interms of Ks Elements. J Aeronaut Aerospace Eng 4: 146. doi:10.4172/2168-9792.1000146

**Copyright:** © 2015 Nair LS. This is an open-access article distributed under the terms of the Creative Commons Attribution License, which permits unrestricted use, distribution, and reproduction in any medium, provided the original author and source are credited.

Citation: Nair LS (2015) Variation in Argument of Perigee for Near-Earth Satellite Orbits Perturbed by Earth's Oblateness and Atmospheric Drag Interms of Ks Elements. J Aeronaut Aerospace Eng 4: 146. doi:10.4172/2168-9792.1000146

$$P_n(x) = \sum_{k=0}^{m} \frac{(-1)^k (2n-2k)! x^{(n-2k)}}{2^n k! (n-k)! (n-2k)!}$$
 where  $m = \frac{n}{2}$  if niseven and  $m = \frac{n-1}{2}$  if nisedd.

As in [16],  $\vec{P} = -\frac{1}{2}\rho \frac{SC_D}{m} |\vec{v}_r| \vec{v}_r$  the effective area of the satellite,

 $C_{\scriptscriptstyle D}$ , the drag coefficient,  $\rho$ , the atmospheric density at the point of calculating atmospheric drag force and  $\vec{v_r}$ , the velocity of the satellite relative to the ambient air. If  $\vec{v}$  is the velocity of the satellite relative to the Earth's centre, then  $\vec{v_r} = \vec{v} - \vec{v_\alpha}$  where  $\vec{v_\alpha}$  is the velocity of the air relative to the Earth's centre.

 $\vec{v}_r \approx \vec{v} \left( 1 - \frac{r_{p_0}}{v_{p_0}} \Lambda \cos i_0 \right)$ ,  $\Lambda$ , the rotational rate of the atmosphere

about the Earth's axis and  $i_o$ , the initial inclination,  $r_{p_0} = a_0(1-e_0)$  is the initial perigee radius,  $\vec{v}_{p_0}$ , the velocity at the initial perigee. Then the drag force per unit mass tangential to the orbit can be written as  $\vec{P} = -\frac{1}{2} \alpha \delta |\vec{v}| \vec{v}$  where  $\delta = \frac{FSC_D}{2}$ , and

$$F = \left(1 - \frac{r_{p_0}}{\vec{v}_{p_0}} \Lambda \cos i_0\right)^2.$$
(6)

Following [17] the density function for an oblateness atmosphere together with day-to-night density variation is

$$\rho = \rho_{p_0} (1 + F \cos \phi) \exp \{-\beta (r - \sigma)\}.$$
(7)

To express  $cos\varphi$  in terms of the true anomaly  $\theta$  and then in terms of the eccentric anomaly E let  $(\alpha_s, \delta_s)$  and  $(\alpha_B, \delta_B)$  are the right ascension and declination of the sun and the day time bulge respectively, then  $\alpha_B = \alpha_s + h; \delta_B = \delta_s$ . We can write

$$\cos\phi = A\cos\theta + B\sin\theta \tag{8}$$

$$A = \sin\delta_{B} \sin i \sin\omega + \cos\delta_{B} \{\cos\Omega - \alpha_{B}\} \sin\omega + \cos i \sin(\Omega - \alpha_{B}) \cos\omega \quad (9)$$

$$B = \sin\delta_B \sin i \sin\omega + \cos\delta_B \left\{ \cos\Omega - \alpha_B \right\} \sin\omega + \cos i \sin(\Omega - \alpha_B) \cos\omega \quad (10)$$

The scale height H is known to increase with altitude and this variation of H will have an influence upon its motion. The value of H may be taken as  $H = H_p + \mu(r - r_p)$  where  $|\mu| < 0.2$  and for any particular value of  $r_p$ . To sum up, expression for the density, similar to that of Swinerd Boulton [17], is

$$\rho_{v} = \left[1 + \frac{1}{2}\mu z^{2}(1 - 2\cos E + \cos^{2} E) + O(\mu c)\right]\rho, \rho \text{ from equation (7) and } z = \frac{ae}{H} \quad (11)$$

### **Analytical Integration**

$$\begin{split} \vec{u} &= (u_1, u_2, u_3, u_4) = \vec{\alpha} cos\left(\frac{E}{2}\right) + \vec{\beta} sin\left(\frac{E}{2}\right) \qquad \vec{u}^* = \frac{d\vec{u}}{dE} = \frac{1}{2} \left[ -\vec{\alpha} sin\left(\frac{E}{2}\right) + \vec{\beta} cos\left(\frac{E}{2}\right) \right] \\ \vec{x} &= (x_1, x_2, x_3) = L(\vec{u})\vec{u}, \qquad r = (x_1^2 + x_2^2 + x_3^2)^{\frac{1}{2}} = u_1^2 + u_2^2 + u_3^2 + u_4^2, \\ \dot{\vec{x}} &= (\dot{\vec{x}}_1, \dot{\vec{x}}_2, \dot{\vec{x}}_3) = L(\vec{u})\vec{u}^* \qquad w = \left[ \frac{1}{2} \left(\frac{K^2}{r} - \frac{1}{2} |\vec{\vec{x}}|^2 - V\right) \right]^{\frac{1}{2}}, \qquad t = \tau - \frac{1}{\omega}(\vec{u}, \vec{u}^*) \\ L(\vec{u}) &= \begin{pmatrix} u_1 & -u_2 & -u_3 & u_4 \\ u_2 & u_1 & -u_4 & -u_3 \\ u_3 & u_4 & u_1 & u_2 \\ u_4 & -u_3 & u_2 & -u_1 \end{pmatrix}. \end{split}$$

In terms of E,

$$r\cos\theta = a(\cos E - e), \quad r\sin\theta = a(1 - e^2)^{\frac{1}{2}}\sin E$$

$$|\vec{v}| = \frac{K^2}{a^{\frac{1}{2}}} \left[ 1 + e\cos E + \frac{e^2}{2}\cos^2 E + \frac{e^3}{2}\cos^3 E \right]$$
(13)

(12)

$$\frac{|\vec{v}|}{r} = \frac{K^2}{a^{\frac{3}{2}}} \left[ 1 + 2e\cos E + \frac{5e^2}{2}\cos^2 E + 3e^3\cos^3 E \right]$$
(14)

The integrals available in the above theory are of the form

$$I(m,n) = \int_0^{2\pi} \cos^m E \, \sin^n E \, dE = 4 \int_0^{\frac{\pi}{2}} \sin^m E \, \cos^n E \, dE$$
, if mand nareeven

and 0 if either m or n is odd.

$$= 2B\left(\frac{m+1}{2}, \frac{n+1}{2}\right) = 2\left(\frac{\Gamma\left(\frac{m+1}{2}\right)\Gamma\left(\frac{n+1}{2}\right)}{\Gamma\left(\frac{m+n+2}{2}\right)}\right), \quad \Gamma\left(\frac{1}{2}\right) = \sqrt{\pi} \text{ and } \Gamma(m) = (m-1)! \text{ if } m > 0$$

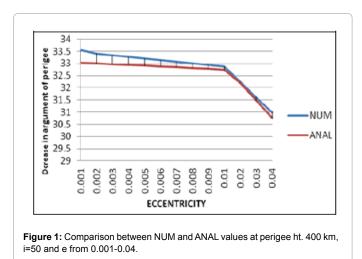
#### **Initial Conditions**

Knowing the position and velocity vectors  $\vec{x}$  and  $\dot{\vec{x}}$  at the instant t = 0, the values of  $r, \omega, t, \vec{u}_i$  and  $\vec{u}_i^*$  can be computed [10], (pp. 91-92), and by adopting E = 0 as the initial value of the eccentric anomaly, we obtain  $\vec{\alpha} = \vec{u}$ ,  $\beta = 2\vec{u}^*$ ,  $\tau = \frac{(\vec{u}, \vec{u}^*)}{w}$ .

# **Numerical Results**

In the entire test cases reported here, the values of  $\omega$ , Right Ascension Node,  $\Omega$  and mean anomaly, M are 60, 30 and 0 degrees respectively. The value of  $K^2$ , R,  $J_2$ ,  $J_3$  and  $J_4$  utilized for numerical computations are  $398600.8km^3s^{-2}$ , 6378.135 kilometer and  $1.0826157 \tilde{A} \times 10^{-3}$ , -2.53648 D - 06 and -1.52 D-06 respectively. Jacchia (1977) atmospheric density model, which is relatively easier to use, is employed to compute the values of  $P_{P_0}$ , the density at the perigee and H, the density scale height at the end of each revolution. Arbitrarily 22 August 2002 is chosen as the initial epoch. The values of  $\varepsilon$ ,  $\wedge$  and  $b_n = (m / c_D A)$ , utilized during the computations are 0.00335, 1.2 and 50.0 respectively. In this model  $c = \frac{1}{2}\beta r_{p_0} sin^2 i$  approaches maximum value 0.2042 at e = 0.003 and  $i = 80^{\circ}$  while minimum value 0.0232 at e = 0.005 and i= 20° respectively at the end of 100 revolutions (Figures 1 and 2). The values of 10.7 cm solar flux ( $F_{10.7}$ ) and averaged geomagnetic index  $(A_p)$  are taken as 150 and 10, respectively, which approximately represents an average density and results in exospheric temperatures between 1000 and 1100 K for the different cases we considered (Table 1).

We have transformed equations for 
$$\frac{d\omega}{dE}, \frac{d\tau}{dE}$$
 and  $\frac{d\tilde{\alpha}_i}{dE}$ , and using



ISSN: 2168-9792 JAAE, an open access journal

Perigee heights in Km								
e	Method	350 Inclination in Degrees			400			
		20	50	80	20	50	80	
0.003	NUM	49.658	33.851	9.213	48.891	33.336	9.069	
	ANAL	49.185	33.585	9.181	48.416	32.971	8.884	
	% ERROR	0.95	0.79	0.35	0.97	1.1	2.04	
0.004	NUM	49.561	33.783	9.193	48.796	33.269	9.05	
	ANAL	49.104	33.584	9.156	48.327	32.942	8.866	
	% ERROR	0.92	0.59	0.4	0.96	0.99	2.04	
0.005	NUM	49.463	33.715	9.173	48.7	33.203	9.032	
	ANAL	49.017	33.582	9.115	48.236	32.912	8.852	
	% ERROR	0.9	0.39	0.64	0.95	0.88	1.99	
0.007	NUM	49.2681	33.58	9.137	48.507	33.071	9	
	ANAL	48.832	33.57	9.063	48.049	32.849	8.833	
	% ERROR	0.89	0.03	0.8	0.94	0.67	1.81	
0.008	NUM	49.171	33.513	9.118	48.411	33.005	8.978	
	ANAL	48.734	33.558	9.039	47.953	32.814	8.825	
	% ERROR	0.89	0.13	0.87	0.95	0.58	1.71	
0.009	NUM	49.073	33.447	9.1	48.316	32.939	8.96	
	ANAL	48.633	33.542	9.024	47.857	32.777	8.817	
	% ERROR	0.9	0.28	0.84	0.945	0.49	1.6	

Table 1: Decrease in Argument of Perigee.

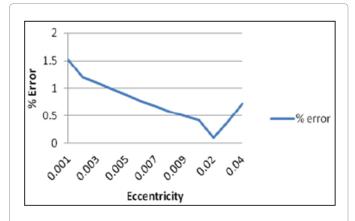
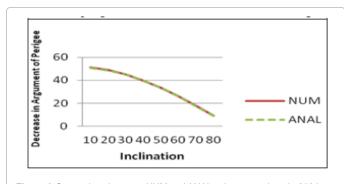


Figure 2: % error between NUM and ANAL values at perigee ht. 400 km, i = 50 and e from 0.001-0.04.



**Figure 3:**Comparison between NUM and ANAL values at perigee ht 350 km, e = 0.01 and i from 10-80 degree.

equations (11), (12), (13), (14) and programmed in double precision arithmetic to compute the KS elements  $\vec{\alpha}_i$  and  $\vec{\beta}_i$  (i = 1,2,3,4). Once

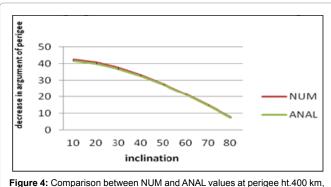


Figure 4: Comparison between NUM and ANAL values at perigee ht.400 km, e=0.1 and i from 10-80 degree.

е	After 1 revolution			After 100		
	NUM	ANAL	% error	NUM	ANAL	% error
0.01	0.5029	0.5019	0.2	48.2202	47.7593	0.96
0.02	0.4931	0.4919	0.24	47.2798	46.7755	1.07
0.03	0.4835	0.4822	0.28	46.3662	45.8329	1.15
0.04	0.4743	0.4728	0.32	45.4785	44.9364	1.19
0.05	0.4653	0.4636	0.36	44.616	44.0643	1.24
0.06	0.4566	0.4548	0.4	43.7776	43.2031	1.31
0.07	0.4481	0.4461	0.44	42.9625	42.3474	1.43
0.08	0.4398	0.4377	0.49	42.1698	41.4952	1.6
0.09	0.4318	0.4295	0.55	41.3988	40.6458	1.82
0.10	0.424	0.4214	0.61	40.6485	39.7981	2.09

 Table 2: Comparison of ANAL to NUM with % error at perigee ht 400 km, i = 80 after 1and 100 revolutions.

 $\vec{\alpha}_i$  and  $\vec{\beta}_i$  are known, they are converted into and  $\vec{x}$  and  $\dot{\vec{x}}$ , which are further converted to the osculating orbital elements (Figures 3 and 4). Percentage Error = (ANAL  $\hat{a} \in \text{NUM}$ )/NUM is calculated to check validity of the work. The algebraic computations are made with MAPLE12 mathematical software (Table 2).

#### Conslusion

The KS element equations are integrated analytically by a series expansion method by assuming an oblate diurnal atmosphere when density scale height varies with altitude and by including the terms corresponds to Earth's zonal harmonics  $J_2$ ,  $J_3$  and  $J_4$ . A wide range of eccentricity and inclination is considered for calculating the change in argument of perigee by present analytical theory and by numerical integrated values for 1 and 100 revolutions shows that the analytically integrated values are reasonably accurate and thus highlights the usefulness of the analytical expressions. Graphical representation as well as the table presented here emphasizes the importance of developing the theory to find the decrease in argument of perigee.

#### References

- De Nike J (1956) The Effect of the Earth's Oblateness and Atmosphere in a Satellite Orbit J of the Franklin Institute Monograph: 79-88.
- Hoots FR (1981)Theory of the Motion of An Artificial Earth Satellite Celestial Mechanics 23: 307-363.
- Lane MH (1965) The Development of an Artificial Satellite Theory Using a Power-Law Atmospheric Density Representation AIAA paper 65-35 AIAA 2 nd Aerospace Sciences Meeting New York.
- 4. Jacchia LG (1964) Static Diffusion Models of the Upper Atmosphere with

Empirical Temperature Profiles SAO SP: 170.

- 5. Jacchia LG (1970)New Static Models of the Thermosphere and Exosphere with Empirical Temperature Profile SAO SP: 313.
- Jacchia LG (1971) Revised Static Models of the Thermosphere and Exosphere with Empirical Temperature Profiles SAO SP: 332.
- Jacchia LG (1977) Thermospheric Temperature Density and Composition: New Models SAO SP: 375.
- Hedin AE (1987) MSIS-86 Thermospheric Model J Geophys Res 92(A5): 4649-4662.
- 9. Hedin AE (1991) Extension of the MSIS Thermosphere Model into the Middle and Lower Atmosphere. J Geophys Res 96(A2): 1159-1172.
- Stiefel EL, Scheifele G (1971) Linear and Regular Celestial Mechanics. Berlin/ Heidelberg/New York 1971. Springer-Verlag.

11. Sharma RK (1993) Analytical Short-Term Orbit Predictions with J3 and J4 in terms of KS Elements. Celes Mech and Dynam Astron 56: 503.

Page 4 of 4

- Sharma RK (1997)Analytical Integration of K-S Elements Equations with J2 for Short-Term Orbit Predictions. Planet Space Sci 45: 1481- 1486.
- Sharma RK (1991) Analytical Approach using KS Elements to Near-Earth Orbit Predictions Including Drag. Proc Roy Soc Lond A 433: 121-130.
- Sharma RK (1997) Contraction of Satellite Orbits using KS Elements in an Oblate Diurnally Varying Atmosphere. Proc Roy Soc Lond A 453: 2353-2368.
- Nair LS, Sharma RK (2003) Decay of Satellite Orbits Using KS Elements in an Oblate Diurnally Varying Atmosphere with Scale-height Dependent on Altitude. Adv Space Res 31: 2011-2017.
- 16. King-Hele DG (1987) Satellite Orbits in an Atmosphere. Theory and Applications Blackie Glasgow and London.
- Swinerd GG, Boulton WJ (1982) Contraction of Satellite Orbits in an Oblate Atmosphere with a Diurnal Density Variation. Proc Roy Soc A 383: 127-145.