



# Temporal Variation of Soil CO<sub>2</sub> Efflux on Sloping Pasture of Heihe River Basin and Effects of Temperature and Soil Moisture

Zongqiang CHANG<sup>1,2\*</sup>, Xiaoqing LIU<sup>1,2</sup>, Qi FENG<sup>1,2</sup> and Xuelong ZHANG<sup>3</sup>

<sup>1</sup>Alashan Desert Eco-hydrology Experimental Research Station, Cold and Arid regions Environmental and Engineering Research Institute, Chinese Academy of Sciences, Lanzhou 730000, China

<sup>2</sup>Gansu Hydrology and Water Resources Engineering Center, Lanzhou 730000, China

<sup>3</sup>Academy of Water Resource Conservation Forests in Qilian Mountains of Gansu Province, Zhangye Gansu 734000, China

## Abstract

Employing LiCor 6400 gas exchange analyzer and soil respiration chamber attachment (LiCor Inc, Lincoln, NE, USA), this paper continuously measured the soil surface CO<sub>2</sub> effluxes on the sloping pasture of Heihe River basin from early April to late October 2010 to investigate the soil CO<sub>2</sub> efflux rate and its feedback to the changes of climate and land use. The results showed that during the growing season, the diurnal variation of pasture soil respiration in the mountain watershed of the Heihe River valley was low at night, with lowest appears at 7:00, 6:30, 5:30, 5:00, 6:00 and 7:00 from May to October, and started to rise rapidly during 7:00~8:30, and then descend during 16:00~18:30. The maximum soil CO<sub>2</sub> efflux appears at 15:00, 14:30, 14:30, 13:30, 14:00 and 15:00. The maximum of average soil CO<sub>2</sub> efflux occurred in July and August, and the second was in May and September, and the third was in April and October. And it was basically consistent in April and October. The diurnal average of pasture soil CO<sub>2</sub> efflux was between 0.31~6.98 μmol m<sup>-2</sup>s<sup>-1</sup>, and the Q<sub>10</sub> value is 2.16. Soil CO<sub>2</sub> efflux had an exponential and Boltzmann correlation with temperature and soil moisture, respectively.

**Keywords:** Soil respiration rate; Growth season; Temperature; Soil moisture; Sloping pasture; Heihe river basin

## Introduction

Soil surface CO<sub>2</sub> efflux is one of the major pathways by which CO<sub>2</sub> fixed by terrestrial plants is released into the atmosphere [1]. Given the controversy over its potential role in amplifying global warming, soil surface CO<sub>2</sub> efflux has recently been the subject of intensive study [2-5]. In the face of impending global warming, increases in soil respiration are likely to mediate progressively lower rates of carbon sequestration [2,6-8]. Despite its obvious importance to carbon cycle processes, soil respiration has proven to be extremely difficult to quantify accurately. Like many other soil processes, respiration exhibits both great spatial heterogeneity, particularly at small spatial scales, and great temporal variability on diurnal, seasonal and inter-annual time scales [9-15].

Sloping pasture is the main pasture types on the mountain watershed of the Heihe River basin, and they cover about 28.27% of the total land area. Understanding of the pasture efflux of CO<sub>2</sub> from the soil surface is a key component of the carbon balance of its ecosystem. Quantifying this flux and understanding the factors that underlie the temperature and soil moisture variation in its magnitude are fundamental to our understanding of the behavior of the ecosystem as a whole and to our ability to predict the likely consequences of climate change [15]. Up to now, a large number of studies have been carried out on the relationship between the dynamics of soil CO<sub>2</sub> flux and related factors [16-19]. However, little has been reported on the sloping pasture soil respiration intensity at regional scale. This study describes the changes in CO<sub>2</sub> flux of sloping pasture soil and the temporal differences during the growing season under different environment conditions were explored in this study. The spatial variations of soil CO<sub>2</sub> flux and its relation to the environmental elements such as soil water content, and temperature were discussed. The results help the scientific community's understanding of carbon exchange mechanisms between soil and atmosphere and the source-sink changes of the terrestrial ecosystem.

## Materials and methods

### Site description

The experimental site is located in the Pailugou watershed in the

Xishui Forest Farm of Sunan County in Gansu Province. The basin covers an area of 2.95 km<sup>2</sup>, with 55% grassland and 40% forest land. The study area has a high, cold, semiarid and sub-humid mountain forest and grassland climate, with mean annual temperature of 0.5°C, mean annual precipitation of 435 mm, and annual potential evaporation of 21051 mm [20]. Sloping pasture mainly occurs in the mountain forest grassland zone at elevation of 2,500~3,000 m and occupies about 28.27% of the Qilian Mountain region's total area. The main plant species are *Carex*, *Achnatherum inebrians*, *Polygonum viviparum*, *Oxytropis ochrocephala*, *Achnatherum splendens*. Soil type in the region is mountain grey cinnamon soil, with a depth of about 1 m. Hence, the study of CO<sub>2</sub> flux of alpine meadow soil at this altitude will contribute to a better understanding of the feedback effect of soil CO<sub>2</sub> emission flux on the climate and land use changes under high altitude and low temperature conditions.

### Soil surface CO<sub>2</sub> efflux measurements

Soil surface CO<sub>2</sub> effluxes were measured on early April to late October 2010 with a LiCor 6400 gas exchange analyzer with soil respiration chamber attachment (LiCor Inc, Lincoln, NE, USA). The 6400 is a closed system infrared gas analyzer that measures soil respiration. We installed five PVC collars (0.008 m<sup>2</sup>) randomly located in each 5m×5m plot 2 days. Collars were imbedded approximately 2 cm into the soil and left in place throughout the measurement period. We did not begin measurements until 1 week after installation to minimize the effects of disturbance from collar installation and removed live vegetation inside

**\*Corresponding author:** Zongqiang Chang, Environmental and Engineering Research Institute, Chinese Academy of Sciences, Lanzhou 730000, China, E-mail: [changzq@lzb.ac.cn](mailto:changzq@lzb.ac.cn)

**Received** January 28, 2013; **Accepted** March 08, 2013; **Published** March 13, 2013

**Citation:** CHANG Z, LIU X, FENG Q, ZHANG X (2013) Temporal Variation of Soil CO<sub>2</sub> Efflux on Sloping Pasture of Heihe River Basin and Effects of Temperature and Soil Moisture. J Geol Geosci 2: 111. doi:10.4172/2381-8719.1000111

**Copyright:** © 2013 CHANG Z, et al. This is an open-access article distributed under the terms of the Creative Commons Attribution License, which permits unrestricted use, distribution, and reproduction in any medium, provided the original author and source are credited.

the chamber collars at least 24 hr before measurements to minimize influences of soil disturbance and root injury on the measurements.

### Soil temperature and moisture measurements

Soil temperature and moisture were determined at each soil CO<sub>2</sub> efflux measurement location and time. Soil moisture within the top 80 mm of soil was measured with a cable tester (1502C Metallic TDR Cable Tester, Tektronix, Inc., Beaverton, OR) and dedicated Time-Domain Reflectometry (TDR) steel probes (8 cm long) inserted vertically from the soil surface. Temperature was measured at a depth of 0.1 m from the soil surface with a copper-constantan thermocouple mounted in an aluminium probe.

### Data analyses

The relationship between soil CO<sub>2</sub> flux and temperature was analyzed using the statistical analysis software SPSS 13.0 for Windows and the dynamical curve was drawn using the Origin Pro 8.0.

## Results

### The diurnal variation in soil CO<sub>2</sub> efflux

Soil CO<sub>2</sub> efflux showed an asymmetric diurnal pattern, with a

minimum between 0300H and 0700H (local time) and a maximum in the early afternoon (13:00H-16:00H). Soil CO<sub>2</sub> efflux followed the increasing trend of soil temperature in the morning, but then leveled off with slight fluctuations, while soil temperature continued to increase in the afternoon. From evening to early morning of the next day, soil CO<sub>2</sub> efflux followed, with few fluctuations, the declining trend of soil temperature. In growth season, Heihe basin mountainous area sloping pasture soil CO<sub>2</sub> efflux daily variation assumes the following characteristic: In the evening maintains at the low level, Minimum value about 6:00, Starts in 7:00~8:30 to elevate, 14:00 about maximizing, 16:00~18:30 to drop gradually, The entire process assumes the single peak curve. Different month, Soil breath speed daily variation existence remarkable difference, displays starts in the soil breath speed to elevate, the maximizing time is different. From May to October, the diurnal variation of soil respiration was low at night, the lowest at 07:00 H, 06:30 H, 05:30 H, 06:00 H and 07:00 H, rise rapidly between 07:00 H and 08:30 H, and then descended between 16:00 H and 18:30 H. The maximum soil CO<sub>2</sub> efflux appeared at 15:00 H, 14:30 H, 14:30 H, 13:30 H, 14:00 H and 15:00 H (Figure 1).

In different periods of growth, soil CO<sub>2</sub> efflux diurnal variation existence remarkable difference (Figure 2) The mean daily soil

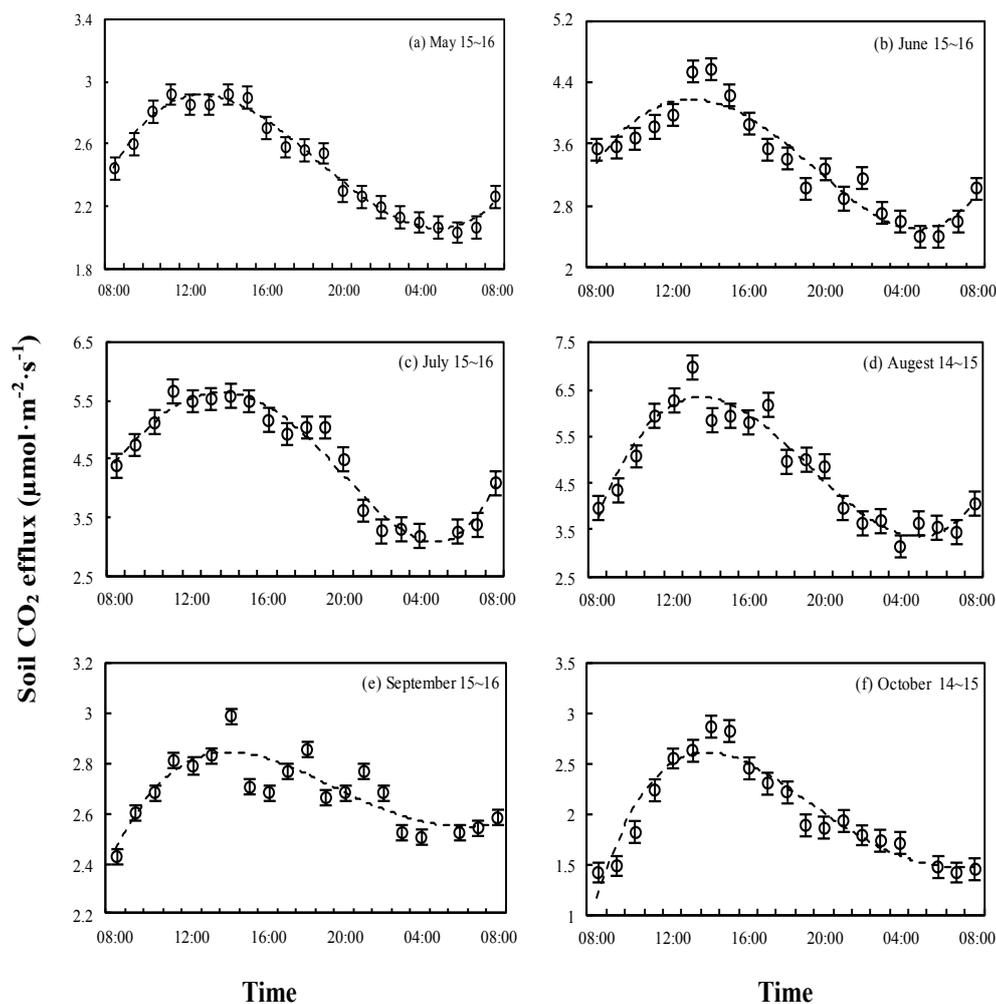


Figure 1: Diurnal variation of meadow soil CO<sub>2</sub> efflux in the mountain watershed of the Heihe River basin.

respiration rate between 0.31~2.58  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$  in May, 0.78~4.85  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$  in June, 4.61~6.98  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$  in July, and then gradually descended at 2.37~6.26  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$  in August, 3.47~4.23  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$  in September, 1.61~4.21  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$  in October.

### The seasonal variation in soil CO<sub>2</sub> efflux

The continuously observed results of CO<sub>2</sub> flux of sloping pasture soil in the Heihe river basin (Figure 3, Table 1) showed that associated with the diurnal variations, soil CO<sub>2</sub> flux was low in the initial growing stage (May) but gradually increased in June, reached a maximum value 8.49  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$  in July to August (Figure 3) and started to decrease in September, the entire process change tendency assumes the single peak curve.

### Effects of soil temperature and moisture on soil CO<sub>2</sub> efflux

Non-linear regression analysis showed that the exponential model

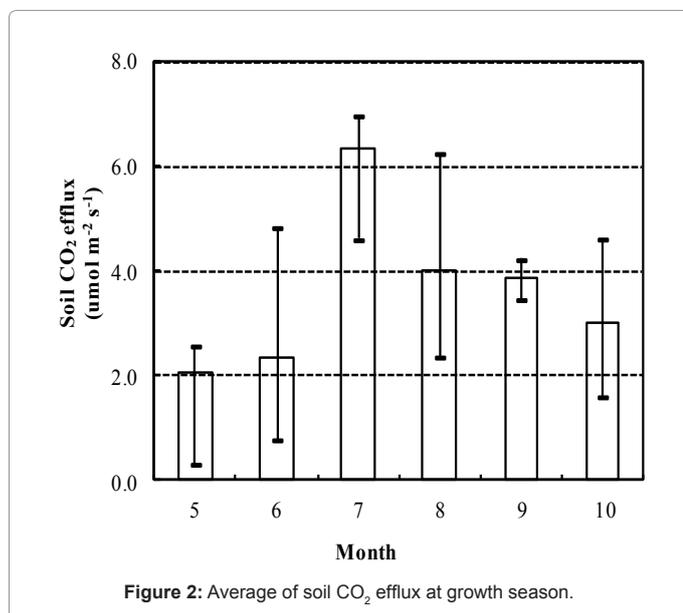


Figure 2: Average of soil CO<sub>2</sub> efflux at growth season.

can better describe the relationship between CO<sub>2</sub> flux of sloping pasture soil in the Heihe river basin and soil temperature at 15 cm depth as shown in figure 4a. Statistical analysis revealed that the correlation between CO<sub>2</sub> flux of alpine meadow soil (S) and soil temperature was significant ( $P < 0.001$ ,  $n = 25$ ), and its regression relation is as follows:

$$S = 1.41e^{0.077t} \quad (R^2 = 0.88)$$

$$Q_{10} = e^{(10 \times 0.077)} = 2.16$$

The variation trends of CO<sub>2</sub> flux of sloping pasture soil in the Heihe river basin and soil water content were not consistent figure 4b. When soil water content was low, the variations in soil CO<sub>2</sub> flux and soil water content were almost synchronous, i.e. CO<sub>2</sub> flux increased with increase in soil water content, but when soil content increased to a certain level the increase in soil CO<sub>2</sub> flux became slow. Analytical results show that the Boltzmann model can better describe the relationship between CO<sub>2</sub> flux of sloping pasture soil in the Heihe river basin and soil moisture ( $P < 0.001$ ,  $n = 25$ ), and its regression relation is as follows:

$$S = 5.54 - 5.03 / (1 + \exp((x - 8.38) / 1.34)) \quad (R^2 = 0.98)$$

### Simultaneous stepwise ranking of variables' effects on soil CO<sub>2</sub> efflux

Soil temperature and soil moisture, as well as their interaction showed effects on changes in soil CO<sub>2</sub> efflux. Using the stepwise regression process in SPSS, all variables were tested simultaneously for their relative contribution to explaining variance in soil CO<sub>2</sub> efflux. These results were similar to those obtained using simple linear regression with individual factors: soil temperature was positively related to soil CO<sub>2</sub> efflux and explained about 46.8% of its variance; soil moisture was weakly positively related to soil CO<sub>2</sub> efflux and explained about 15.7% of its variance; All of the variables combined explained about 62% of the variation in annual CO<sub>2</sub> efflux in 2010.

### Discussion

Soil CO<sub>2</sub> efflux showed diurnal and seasonal changes (Figure 1,2) The measurements of soil CO<sub>2</sub> efflux presented here for the sloping pasture sites were in the range of 0.31-6.98  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ . This result is

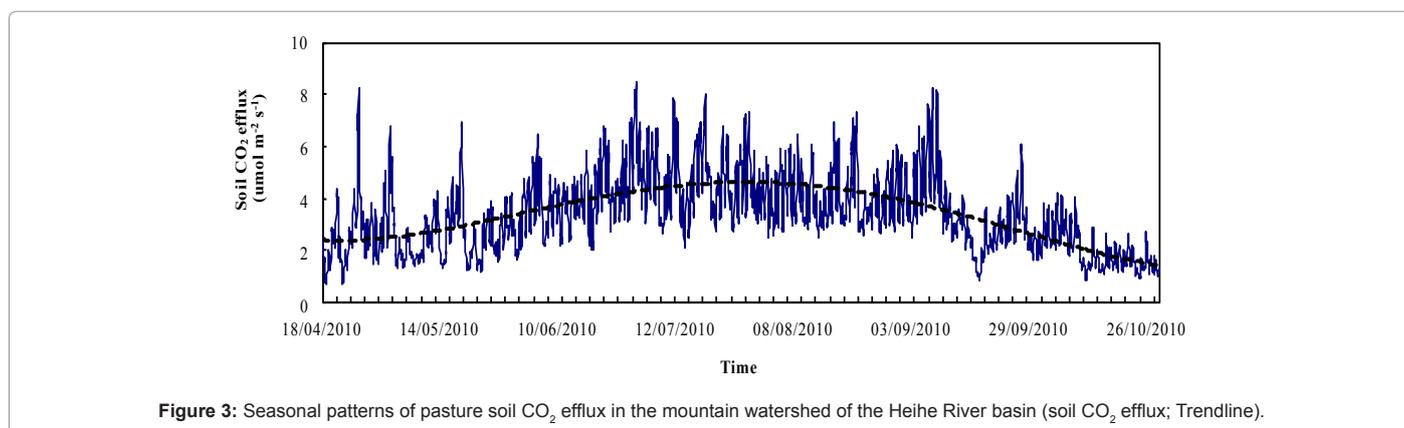
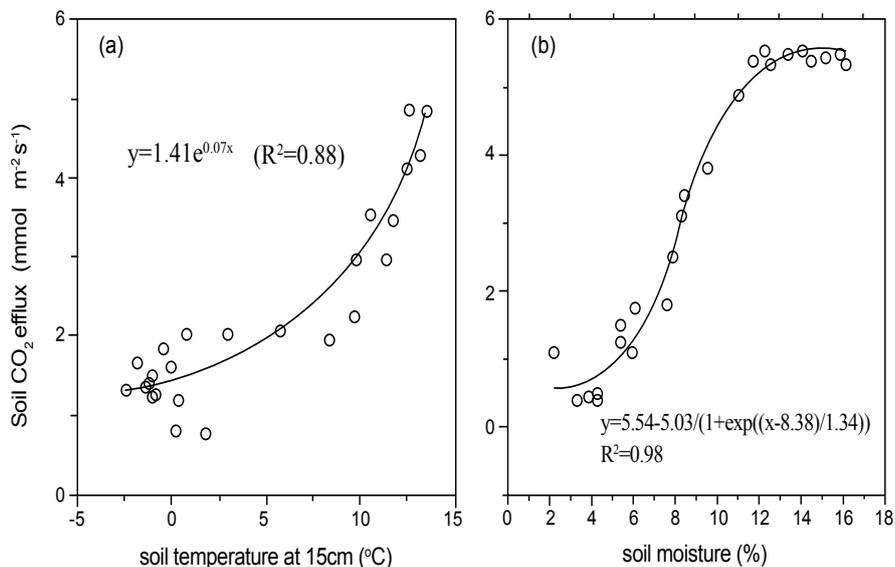


Figure 3: Seasonal patterns of pasture soil CO<sub>2</sub> efflux in the mountain watershed of the Heihe River basin (soil CO<sub>2</sub> efflux; Trendline).

Month	5	6	7	8	9	10
Mean Soil CO <sub>2</sub> efflux/ $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$	2.06	2.37	6.35	3.99	3.85	3.01
maximum value/ $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$	6.35	8.43	8.49	8.37	8.24	4.21
minimum value/ $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$	1.03	1.54	2.11	2.48	0.88	0.82
CV / %	63.4	56.0	52.4	55.3	61.6	54.6

Table 1: Means and coefficients of variance of pasture soil CO<sub>2</sub> efflux.



**Figure 4:** The relationship between soil respiration rate, soil temperature at 15cm depth (a) and soil moisture (b) on pasture in the mountain watershed of the Heihe River basin.

higher than one of 2.32-5.70  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$  reported by Michael [21] for grassland in northern Ontario of Canada. Raich and Schlessinger [6] summarizes large amounts of data showed that average  $Q_{10}$  of grassland around 2.4, but for forest soils, Liu Shaohui and Fang jinyun [22] believes that global average  $Q_{10}$  is 1.57. This study indicated that the  $Q_{10}$  value of sloping pasture soil is 2.16, And LI Linhao [23] in sheep grass grassland research results showed that  $Q_{10}$  value is 2.0~3.0 (to temperatures for based on), slightly than tropical dilute tree grassland in North Australia [24] and high grass grassland in North America [25] they think this phenomenon attributed to research locations of latitude location partial high, because  $Q_{10}$  value in cold region higher than warm region [26]. This study site and the leymus chinensis grassland study sites at the same temperate zone, and temperature ranges of the experimental period commensurate with each other so their  $Q_{10}$  values are closer.

While multiple factors contribute to the differences in measured efflux rates, the generally low soil moisture and high soil temperature at our site are likely to be the two major factors which determined the magnitude of the soil CO<sub>2</sub> efflux at the Heihe river basin. Jensen et al. [27] measured soil-surface CO<sub>2</sub> efflux over two days at 8 locations in a *Pinus radiata* D. Don forest in New Zealand using a dynamic chamber method (portable infra-red CO<sub>2</sub> analyser). Their results showed no apparent diurnal pattern in CO<sub>2</sub> efflux, which may have been a result of a lack in variation in soil temperature (at 150 mm depth) and the high soil moisture (close to field capacity during the measurements). However, Davidson et al. [28] reported a diurnal trend resembling the temperature pattern. Kutsch & Kappen's [29] measurements in crop fields showed a diurnal trend of CO<sub>2</sub> efflux similar to ours, except that their diurnal maximum occurred later (about 16:00 H).

When simultaneously considered, soil temperature and soil moisture explained 46.8 and 15.7, respectively, of the variance in soil CO<sub>2</sub> efflux on our sites. Thus our results indicate that soil temperature as a single factor explains the greatest amount of variance in soil CO<sub>2</sub> efflux observed within and across sites and over seasons in the Heihe river basin ( $R^2=0.88$ ,  $P<0.0001$ ). Our findings are consistent with

reports that cite a strong relationship between soil temperature and soil CO<sub>2</sub> efflux [30,31].

Generally, soil moisture limits soil CO<sub>2</sub> efflux at either extremely high or low moisture levels [30-32]. In agreement with the results of Davidson et al. [28], at the Heihe river basin the single factor of soil moisture was correlated ( $R^2=0.98$ ,  $P<0.0001$ ,  $N=25$ ) with CO<sub>2</sub> efflux, and the soil CO<sub>2</sub> efflux increased slow at high soil moisture contents (>15%). This effect at high soil moisture may also be related to the availability of O<sub>2</sub> in the soil pore space, which affects microbial activity. From laboratory and theoretical studies some researchers have found that high water content can impede diffusion of O<sub>2</sub> into the soil, which in turn impedes decomposition and CO<sub>2</sub> production [33-34].

## Conclusions

Soil-surface CO<sub>2</sub> efflux measurements were made on sloping pasture of Heihe river basin from April to October 2010. During the growing season, the diurnal variation of pasture soil respiration in the mountain watershed of the Heihe River valley was low at night, with lowest appears at 7:00, 6:30, 5:30, 5:00, 6:00 and 7:00 from May to October, and started to rise rapidly during 7:00~8:30, and then descend during 16:00~18:30. The maximum soil CO<sub>2</sub> efflux appears at 15:00, 14:30, 14:30, 13:30, 14:00 and 15:00. The maximum of average soil CO<sub>2</sub> efflux occurred in July and August, and the second was in May and September, and the third was in April and October and it was basically consistent in April and October. The diurnal mean soil-surface CO<sub>2</sub> efflux of sloping pasture stand ranged from 0.31  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$  to 6.98  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ , and the  $Q_{10}$  value is 2.16. The positive relationship between soil CO<sub>2</sub> efflux and soil temperature that was observed is well documented. And the soil CO<sub>2</sub> efflux showed a Boltzmann correlation with soil water content.

## Acknowledgements

This work was funded by the National Natural Science Foundation of China (91025002; 30970492; 31270482; 41271037), National Natural Science Foundation of Gansu Province (1107RJZA089) and the National Key Technology R&D Program (2012BAC08B05). The authors would like to express their deep

gratitude to the anonymous reviewers for their valuable suggestions that greatly improved the manuscript.

## References

- Sanchez ML, Ozores MI, Lopez MJ, Colle B, Torre De, et al. (2003) Soil CO<sub>2</sub> fluxes beneath barley on the central Spanish plateau. *Agricultural and Forest Meteorology* 118: 85-95.
- Trumbore SE, Chadwick OA, Amundson R (1996) Rapid exchange between soil carbon and atmospheric carbon dioxide driven by temperature change. *Science* 272: 393-396.
- Liski J, Ilvesniemi H, Maklela A, Westman CJ (1999) CO Emissions from Soil in Response to Climatic Warming Are Overestimated: The Decomposition of Old Soil Organic Matter Is Tolerant of Temperature. *Ambio* 28: 171-174.
- Cox PM, Betts RA, Jones CD, Spall SA, Totterdell IJ (2000) Acceleration of global warming due to carbon-cycle feedbacks in a coupled climate model. *Nature* 408: 184-187.
- Luo Y, Wan S, Hui D, Wallace LL (2001) Acclimatization of soil respiration to warming in a tall grass prairie. *Nature* 413: 622-625.
- Raich JW, Schlesinger WH (1992) The global carbon dioxide flux in soil respiration and its relationship to vegetation and climate. *Tellus* 44B: 81-89.
- Woodwell GM, Mackenzie FT, Houghton RA, Apps M, Gorham E, et al. (1998) Biotic feedbacks in the warming of the earth. *Climate Change* 40: 495-518.
- Davidson EA, Verchot LV, Cattanio JH, Ackerman IL, Carvalho JEM (2000) Effects of soil water content on soil respiration in forests and cattle pastures of eastern Amazonia. *Biogeochemistry* 48: 53-69.
- Zogg GP, Zak DR, Burton AJ, Pregitzer KS (1996) Fine root respiration in northern hardwood forests in relation to temperature and nitrogen availability. *Tree Physiology* 16: 719-725.
- Law BE, Baldocchi DD, Anthoni PM (1999) Below-canopy and soil CO<sub>2</sub> fluxes in a ponderosa pine forest. *Agricultural and Forest Meteorology* 94:171-188.
- Stoyan H, De-Polli H, Bohm S, Robertson GP, Paul EA, et al. (2000) Spatial heterogeneity of soil respiration and related properties at the plant scale. *Plant and soil* 222: 203-214.
- Buchmann N (2000) Biotic and abiotic factors controlling soil respiration rates in *Picea abies* stands. *Soil Biology & Biochemistry* 32: 1625-1635.
- Casals P, Romanya J, Cortina J, Botner P, Couteaux MM, et al. (2000) CO<sub>2</sub> efflux from a Mediterranean semi-arid forest soil. I. Seasonality and effects of stoniness. *Biogeochemistry* 48: 261-281.
- Savage KE, Davidson EA (2001) Interannual variation of soil respiration in two New England forests. *Global Biogeochemical Cycles* 15: 337-350.
- Xu M, Qi Y (2001) Spatial and seasonal variations of Q<sub>10</sub> determined by soil respiration measurements at a Sierra Nevada forest. *Global Biogeochemical Cycles* 15: 687-696.
- Shi WY, Zhang JG, Yan MJ, Yamanaka N, Sheng Du (2012) Seasonal and diurnal dynamics of soil respiration fluxes in two typical forests on the semiarid Loess Plateau of China: Temperature sensitivities of autotrophs and heterotrophs and analyses of integrated driving factors. *Soil Biology and Biochemistry* 52: 99-107.
- Bingrui Jia, Guangsheng Zhou (2009) Integrated diurnal soil respiration model during growing season of a typical temperate steppe: Effects of temperature, soil water content and biomass production. *Soil Biology and Biochemistry* 41: 681-686.
- Epron D, Farque L, Lucot E, Badot PM (1999) Soil CO<sub>2</sub> efflux in a beech forest: dependence on soil temperature and soil water content. *Annals of Forest Science* 56: 221-226.
- Susfalk RB, Cheng WX, Johnson DW, Walker RF, Verburg P, et al. (2002) Lateral diffusion and atmospheric CO<sub>2</sub> mixing compromise estimates of rhizosphere respiration in a forest soil. *Can J For Res* 32: 1005-1015.
- Chang Z, Feng Qi, Jianhua Si, Yonghong Su, Haiyang Xi, et al. (2009) Analysis of the spatial and temporal changes in soil CO<sub>2</sub> flux in alpine meadow of Qilian Mountain. *Environmental Geology* 58: 483-490.
- Laporte MF, Duchesne LC, Wetzel S (2002) Effect of rainfall patterns on soil surface CO<sub>2</sub> efflux, soil moisture, soil temperature and plant growth in a grassland ecosystem of northern Ontario, Canada: implications for climate change. *BMC Ecology* 2: 10.
- Liu shaohui, Fang Jingyun (1997) Effect factors of soil respiration and the temperature's effects on soil respiration in the global scale. *Acta ecologica snica* 17: 469-476.
- Linhao Li, Qibing W, Yongfei Bai, et al. (2000) Soil respiration of a *Leymus* Chinese grassland stand in the Xilin River basin as affected by over-grazing and climate. *Acta Phytocologica Sinica* 24: 680-686.
- Bridge NK, Mott JJ, Hartigan RJ (1983) The formation of degraded area in the dry savanna woodlands of northern Australia. *Australian Journal of Soil Research* 21: 91-104.
- Kucera C, Kirkham D (1971) Soil respirations studies in tall grass prairie in Missouri. *Ecology* 5: 912-915.
- Kirschbaum MUF (1995) The temperature dependence of soil organic matter decomposition, and the effect of global warming on soil organic C storage. *Soil Biology and Biochemistry* 27: 753-760.
- Jensen LS, Mueller T, Tate KR, Ross DJ, Magid J, et al. (1996) Soil surface CO<sub>2</sub> flux as an index of soil respiration in situ: a comparison of two chamber methods. *Soil Biology and Biochemistry* 28: 1297-1306.
- Davidson EA, Belk E, Boone RD (1998) Soil water content and temperature as independent or confounded factors controlling soil respiration in a temperate mixed hardwood forest. *Global Change Biology* 4: 217-227.
- Kutsch WL, Kappen L (1997) Aspects of carbon and nitrogen cycling in soils of the Bornhöved Lake district II, Modeling the influence of temperature increase on soil respiration and organic carbon content in arable soils under different managements. *Biogeochemistry* 39: 207-224.
- Kowalenko CG, Iverson KC, Cameron DR (1978) Effect of moisture content, temperature and nitrogen fertilization on carbon dioxide evolution from field soils. *Soil Biol Biochem* 10: 417-423.
- Pangle RE, Seiler J (2002) Influence of seedling roots, environmental factors and soil characteristics on soil CO<sub>2</sub> efflux rates in a 2-year-old loblolly pine (*Pinus taeda* L.) plantation in the Virginia Piedmont. *Environ Pollut* 116: S85-S96.
- Howard DM, Howard PJA (1993) Relationships between CO<sub>2</sub> evolution, moisture content, and temperature for a range of soil types. *Soil Biol Biochem* 25: 1537-1546.
- Linn DM, Doran JW (1984) Effect of water-filled pore space on carbon dioxide and nitrous oxide production in tilled and non-tilled soils. *Soil Science Society of American Journal* 48: 1267-1272.
- Doran JW, Mielke IN, Power JF (1991) Microbial activity as regulated by soil water-filled pore space. In: *Trans. 14th International Congress of Soil Science. International Society of Soil Science Kyoto Japan* 94-99.