

Radioactive Disequilibrium Studies in Uranium Series of Core Samples of Koppunuru Area, Guntur District, Andhra Pradesh, India

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Abstract

Disequilibrium studies were attempted on mineralised core samples (n=870) from Koppunuru uranium deposit located in south-western part of Palnad sub-basin, Guntur district, Andhra Pradesh, India. The area exposes Banganapalle quartzites unconformably deposited over altered biotite granite (basement). Uranium mineralisation in Koppunuru deposit is hosted by Banganapalle quartzites well above the unconformity, and grit/basement granite close to the unconformity contact. For disequilibrium studies, the core samples were broadly divided in two groups, (a) quartzite hosted (above unconformity) and (b) basement granite hosted mineralisation (below unconformity). Average disequilibrium factor of 41% has been recorded in favour of parent uranium in both types of core samples. It shows significant enrichment of uranium in the system as evident from 41% of disequilibrium in favour of parent uranium. This is probably due to significant migration of some of the daughter radio nuclides due to dissolution of minerals by groundwater action. Besides, the escape of radioactive radon might have accentuated the disequilibrium factor thus increasing the grade of uranium mineralization. The presence of fractures and faults in the study area are the probable conduits for radon migration/escape. Linear regression coefficient between uranium and radium is 0.98 indicates invariability of disequilibrium irrespective of grade.

Keywords: Uranium; Disequilibrium factor; Beta-gamma method; Gamma ray spectrometry; Guntur district; Andhra Pradesh

Introduction

Linkage between economic growth of any country and energy requirement is well known, and hence, sustainable energy resources are essential. During the past three decades the world was able to cope with an increasing energy demand by relying more on fossil fuel [1-3]. However, the progressively dwindling reserves of fossil fuel and a deep-rooted concern about global warming, arising out of CO₂ emission due to the excessive use of fossil fuel, have now developed a growing interest in nuclear energy as an alternative green source [4,5]. Uranium is one of the main nuclear fuels, continuous supply of which is needed for sustainable development in the energy field. In India, an extensive exploration programme is being carried out in different geological domains to establish new uranium resources and reserves to overcome the demand and supply gap in nuclear energy sector. The uranium series disequilibrium in radioactive ore poses a critical problem for proper assessment of the resource in most of the cases. However, it is observed that the magnitude and frequency of radioactive disequilibria is generally ignored which leads to underestimation or overestimation of ore reserve [6,7]. Thus, to overcome uranium ore deposit evaluation related constraints, the disequilibrium studies are significant both in field and laboratory counting measurements.

In the uranium series, the system is considered to be in radioactive equilibrium when all the daughter products decay at the same rate in which they are produced from the parent isotope in an ideally closed system [6,8,9]. Thus, at radioactive equilibrium, each of the daughter products would be present in a constant proportion to the parent isotope. However, such equilibrium is rarely observed in uranium ore bodies [7,9]. Various studies have been undertaken on the nature and significance of disequilibrium conditions in uraniumiferous deposits [10], which have indicated that different geologic and physicochemical processes influence the system. The addition or removal of any isotope in the disintegration series in a radioactive mineral causes disequilibrium in the proportions of the parent isotope to its daughter products [9]. The present paper deals with the evaluation of disequilibrium pattern

in unconformity proximal and fracture controlled types of uranium mineralisation in Koppunuru area, Guntur district, Andhra Pradesh, India using beta-gamma and gamma-ray spectrometry techniques on mineralised borehole core samples.

Geological Setting

Koppunuru area is located in the south-western part of the Neoproterozoic Palnad Sub-basin, where Kurnool sediments are deposited unconformably over the Archaean to Palaeoproterozoic basement granite and gneisses [11,12]. Basement granites are exposed as an inlier over an area of 5 km x 2.5 km to the south-east of Koppunuru, and along the up-thrown block of regional WNW-ESE trending fault to the south of Koppunuru (Figure 1). These are dissected by a number of ENE-WSW, NE-SW, NW-SE and a few N-S trending lineaments represented by dolerite dykes, fractures and faults [13]. The Kurnool sediments are mainly represented by Banganapalle, Narji and Paniam formations comprising quartzites, shales and limestones. The Banganapalle Formation (10-173m thick), the oldest sedimentary lithounit in the study area, and comprises quartzite and intercalated grey shale sequence with basal conglomerate/grit [12].

Proterozoic unconformity-related uranium deposit in Koppunuru and adjoining areas is predominantly (~85%) hosted by Banganapalle quartzite and grit. At places the mineralisation also transgresses below

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Received December 15, 2016; Accepted January 19, 2017; Published January 25, 2017

Citation: Srinivas Y, Singh RV, Rahul Banerjee, Sharma PK, Verma MB (2017) Radioactive Disequilibrium Studies in Uranium Series of Core Samples of Koppunuru Area, Guntur District, Andhra Pradesh, India. J Geol Geophys 6: 277. doi: 10.4172/2381-8719.1000277

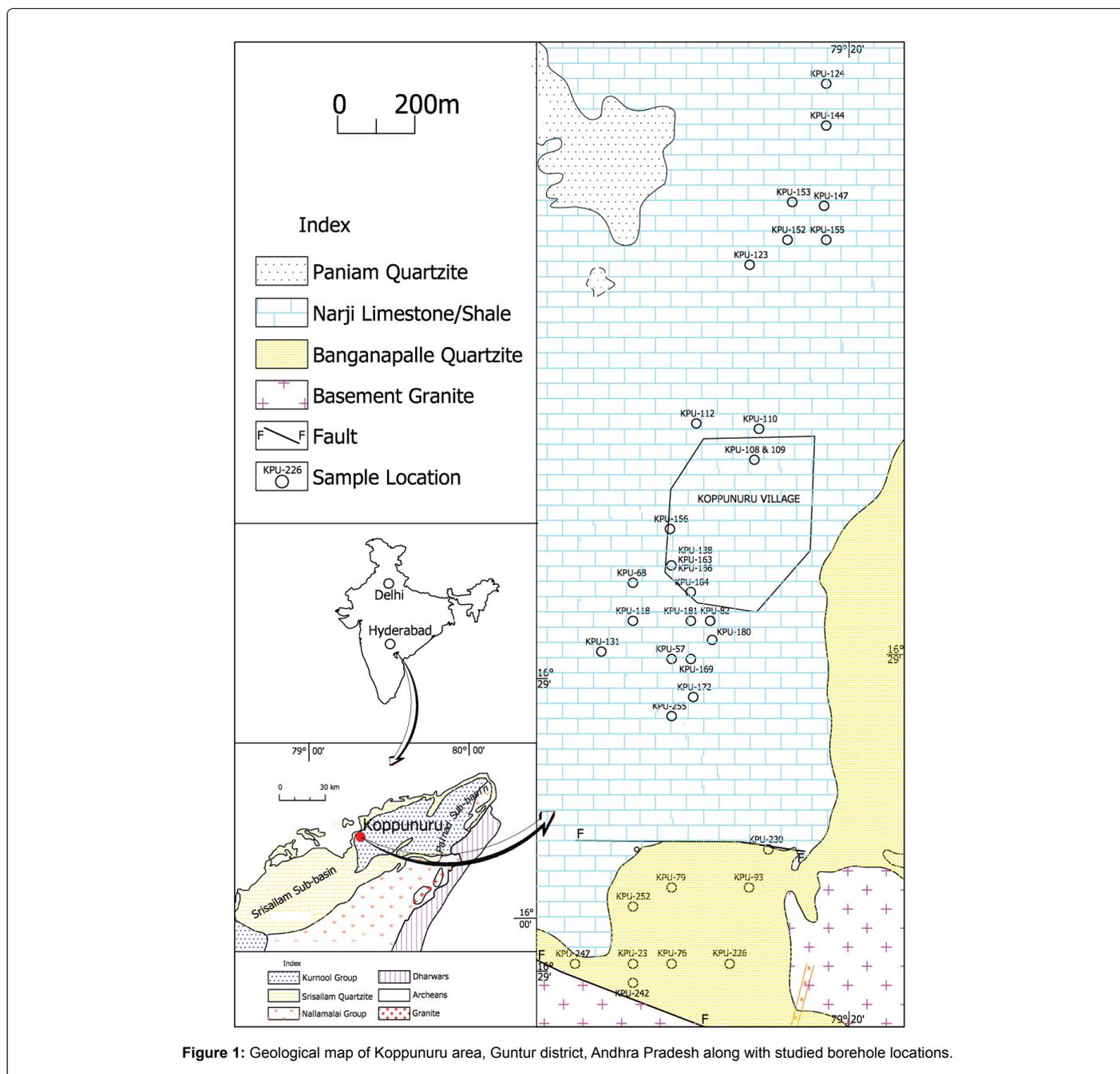
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the unconformity contact in basement granites along the fracture planes [7,14-19]. It is observed that the mineralisation follows a predominant N-S to NNE-SSW trend sympathetic to faults and fractures in the area. Pitchblende and coffinite are identified as primary uranium ore minerals in Koppunuru deposit, while uranophane, phosphuranylite, metazeunerite and U-Ti complex occur as secondary uranium minerals [15,17].

Sampling and Analytical Techniques

A total of 870 mineralised core samples from 34 boreholes of Koppunuru area have been collected for disequilibrium studies, which includes 620 quartzite/grit and 250 granite samples. Drilling was

done using different capacity mechanical coring rigs (Rock Drill-30, Drill Max-400 and Trolley mounted Rock Drill-300). Uranium mineralisation has been intercepted along the studied boreholes at different depths ranging from 26.0 m to 170.0 m. Different mineralised bands show grade and thickness ranging from 0.01% to 0.322% eU_3O_8 and 0.6 m to 7.0 m respectively. Details of mineralized intercepts, lithounit, depth of unconformity, and number of samples collected is shown in Table 1. The mineralised core samples were crushed to -200 mesh to homogenize. 50 g and 140 g of sample were taken after coning and quartering for determination of U_3O_8 (%) by β/γ method, and equivalent U_3O_8 (% eU_3O_8), radium equivalent U_3O_8 (%Ra eU_3O_8), % ThO_2 and % K using High Energy Gamma Ray Spectrometry (HEGS) respectively [6,20]. The representative samples were transferred to



S.No	BH NO.	Mineralised zone		Thickness x Ave.Grade (%eU ₃ O ₈)	Rock Type	Depth of unconformity (m)	n
		from (m)	to (m)				
1	KPU/23	101.25	101.85	0.60 m of 0.040	Gritty Quartzite	107	8
		104.75	106.45	1.70 m of 0.090	Gritty Quartzite		
2	KPU/57	106.21	107.21	1.00 m of 0.319	Granite	104.3	9
3	KPU/68	53.15	53.75	0.60 m of 0.023	Gritty Quartzite	56	3
		55.75	56.75	1.00 m of 0.031	Granite		2
4	KPU/76	63.15	64.35	1.20 m of 0.088	Gritty Quartzite	65.2	7
5	KPU/79	54.88	57.95	3.07 m of 0.130	Gritty Quartzite	58.5	23
		60.05	61.27	1.22 m of 0.130	Granite		10
6	KPU/82	160	161.4	1.40 m of 0.041	Gritty Quartzite	162	3
		163.8	164.8	1.00 m of 0.016	Granite		2
7	KPU/93	104.65	108.45	3.80 m of 0.130	Quartzite/Shale	119	11
		114.95	116.35	1.40 m of 0.017	Gritty Quartzite		
8	KPU/108	156.65	157.65	1.00 m of 0.043	Gritty Quartzite	161	5
		158.95	159.25	0.30 m of 0.011	Gritty Quartzite		
9	KPU/110	140.85	142.05	1.20 m of 0.021	Granite	140.7	14
		148.63	149.79	1.16 m of 0.031	Granite		
10	KPU/112	165.25	168.85	3.60 m of 0.051	Quartzite/Shale	169.45	28
					Gritty Quartzite		
11	KPU/118	140.85	142.45	1.60 m of 0.100	Gritty Quartzite	145.15	22
		145	146	1.00 m of 0.019	Granite		9
12	KPU/123	159.1	163.1	4.00 m of 0.028	Gritty Quartzite	167.4	17
13	KPU/124	125.96	127.1	1.14 m of 0.224	Gritty Quartzite	127.8	16
		129.9	131.38	1.48 m of 0.020	Granite		24
14	KPU/131	147.24	148.43	1.19 m of 0.017	Gritty Quartzite	149.6	23
		148.72	150.03	1.31 m of 0.016	Granite		4
15	KPU/138	141.64	147.08	5.44 m of 0.322	Gritty Quartzite	156.4	93
		147.9	152.81	4.91 m of 0.098	Gritty Quartzite		
		154.14	155.29	1.15 m of 0.025	Gritty Quartzite		
		155.4	158.11	2.67 m of 0.020	Granite		
		160.38	161.65	1.27 m of 0.018	Granite		
16	KPU/144	128.38	130.04	1.66 m of 0.175	Gritty Quartzite	134.15	12
		135.22	136.26	1.04 m of 0.024	Granite		
		139.04	140.27	1.23 m of 0.021	Granite		
		149.15	150.61	1.46 m of 0.032	Granite		
17	KPU/147	128.05	129.25	1.20 m of 0.090	Gritty Quartzite	142.8	3
		143.59	144.88	1.29 m of 0.022	Granite		13
18	KPU/152	135.95	138.05	2.10 m of 0.270	Gritty Quartzite	155.7	21
		155.8	156.4	0.60 m of 0.026	Granite		
		160.75	161.8	1.05 m of 0.031	Granite		10
19	KPU/153	113	114.07	1.07 m of 0.019	Gritty Quartzite	151.2	13
		132.75	133.75	1.00 m of 0.073	Gritty Quartzite		
20	KPU/155	121.5	122.7	1.30 m of 0.074	Gritty Quartzite	138.5	6
21	KPU/156	154.4	155.75	1.35 m of 0.017	Gritty Quartzite	162.35	16
22	KPU/163	138.31	139.34	1.03 m of 0.067	Gritty Quartzite	143.8	28
		142.15	143.66	1.51 m of 0.099	Gritty Quartzite		
		154.86	157	2.14 m of 0.152	Granite		37
23	KPU/166	146.22	147.42	1.20 m of 0.027	Gritty Quartzite	156.9	22
		155.05	156.85	1.80 m of 0.043	Gritty Quartzite		
		156.9	159.3	2.40 m of 0.027	Granite		20
24	KPU/169	90.45	91.65	1.20 m of 0.048	Gritty Quartzite	98.8	12
25	KPU/172	40.03	41.11	1.08 m of 0.081	Quartzite	84.3	20
		80.15	81.85	1.70 m of 0.038	Gritty Quartzite		
		86.65	86.55	0.90 m of 0.012	Granite		6
26	KPU/180	37.41	39.88	2.47 m of 0.139	Quartzite	91	28
		79.38	80.42	1.04 m of 0.101	Gritty Quartzite		
		81.05	82.25	1.20 m of 0.067	Gritty Quartzite		

27	KPU/181	106.55	108.55	1.20 m of 0.065	Gritty Quartzite	108.8	16
		108.9	109.2	0.30 m of 0.200	Granite		1
28	KPU/184	115.65	116.67	1.02 m of 0.018	Gritty Shale	128.8	52
		121.55	128.55	7.00 m of 0.038	Gritty Quartzite		14
		139.55	141.45	1.90 m of 0.025	Granite		
29	KPU/226	38.35	39.65	1.30 m of 0.031	Gritty Quartzite/Shale	40.65	7
30	KPU/230	25.45	27.35	1.90 m of 0.057	Conglomerate	27.9	10
31	KPU/242	79.05	80.95	1.90 m of 0.213	Shale/Quartzite	87.7	13
32	KPU/247	86.48	88.95	2.47 m of 0.105	Quartzite	111.55	45
		104.2	106.6	2.50 m of 0.017	Grit/Conglomerate		
		109.15	110.45	1.30 m of 0.020	Grit/Conglomerate		
		112.15	112.75	0.60 m of 0.016	Granite		10
		119	119.6	0.60 m of 0.023	Granite		
33	KPU/252	69.1	70.6	1.60 m of 0.020	Quartzite	73.4	9
		73.5	75.7	2.20 m of 0.018	Granite		7
		84.3	85.3	1.00 m of 0.011	Granite		
34	KPU/255	88.08	89.19	1.11 m of 0.039	Quartzite	100.9	28
		94.35	96.25	1.90 m of 0.031	Conglomerate		
Total							870

Table 1: Details of mineralised zones in boreholes and host rock of Koppunuru area, Guntur district, Andhra Pradesh.

airtight plastic containers and kept for about a month for attainment of radioactive equilibrium between radon daughters and parent radium in the uranium series [21,22]. Besides, U_3O_8 contents in the samples were also estimated by simultaneous measurement of total beta and total gamma radiations using a LND 73201 beta tube, and 1.75" x 2" NaI(Tl) scintillation detector, respectively in the samples [23,24].

Estimation of uranium

The concentration of U_3O_8 in the sample was estimated by simultaneous measurement of total beta and total gamma radiations by beta gamma method using equation:

$$U_3O_8 = (1+C)U_{\beta} - CU_{\gamma} \quad [25] \quad (1)$$

Where U_{β} = activity of uranium in sample.

U_{γ} = activity of uranium in sample.

C = ratio of Ra_{β} to U_{β} in standard.

The detection limit is 90 ppm with $\pm 10\%$ error. For accuracy an IAEA reference standard RGU-1 (U_3O_8 value 460 ppm and $Ra(eU_3O_8)$ value 470 ppm) was also analysed (n=5). The U_3O_8 value obtained for RGU-1 by this method was 453 ± 24 ppm.

Estimation of $Ra(eU_3O_8)$, ThO_2 and K

$Ra(eU_3O_8)$, ThO_2 and K concentrations in the samples were estimated by using gamma ray spectrometry. A 5" x 4" NaI(Tl) scintillation detector was used for the analysis. The detector was coupled to a dMCA-pro-digital- Multi Channel Analyser (Terjet, Germany). The dMCA directly digitizes signals from the radiation detector and stores them in the format desired by the inbuilt software (winTMCA32). For the estimation of $Ra(eU_3O_8)$, the 1.76 MeV gamma ray energy was measured from the Bi-214 as the daughter of radium series always remains in equilibrium with radium. The estimation of ThO_2 was done by measuring the 2.62 MeV gamma ray energy from Tl-208 and the 1.46 MeV of gamma ray energy was measured for the estimation of % K. Prior to this, energy calibration was done using standard gamma ray sources ^{137}Cs 662 KeV and ^{60}Co 1173KeV and 1332 KeV energies. The stripping and sensitivity factors were calculated using standard reference material RGU-1, RGTh-1 and RGK-1 supplied by IAEA,

Vienna. An In-house (developed at Atomic Minerals Directorate for Exploration and Research, Hyderabad) equilibrium U_3O_8 standard was also used for sensitivity calculations. The samples and standards were taken in the plastic containers of the same volume and size to maintain a same counting geometry to minimize the geometrical error. The containers were sealed carefully to avoid the escape of radon gas from the samples. The counting of samples was done in a Low Background room which is 4 ft below the ground level and walls of the room are made of quartz, with a thickness of 0.9 m. The energy spectra from each sample were obtained by placing the sample on the top of detector.

The $Ra(eU_3O_8)$ concentration was calculated by dividing the net peak area of the characteristic gamma ray energy of 1.76 MeV to the sensitivity of radium [26]. Sensitivity of $Ra(eU_3O_8)$ with a counting time of 200 s is 6.5 counts/ppm for 140 g of sample weight and detection limit is 2 ppm (error <10%). ThO_2 concentration was calculated by dividing the net peak area of the characteristic gamma ray energy of 2.62 MeV to the sensitivity of thorium and similarly the concentration of % K is calculated by dividing the net peak area of characteristic gamma ray energy of 1.46 MeV to the sensitivity of potassium. Sensitivity of ThO_2 with a counting time of 200 s is 2.1 counts/ppm for 140 g of sample weight and detection limit is 5 ppm (error <10%). The net peak area of gamma ray was obtained by subtracting background net counts and stripping of the higher energy contribution.

Results and Discussion

The uranium concentration, $Ra(eU_3O_8)$ concentration of 870 core samples from 34 boreholes of Koppunuru deposit with disequilibrium factor (DF), number of samples of both above and below unconformity, mineralised depth ranges, and minimum, maximum and average values of U_3O_8 , $Ra(eU_3O_8)$ and ThO_2 of the studied core samples are given in Table 2.

The linear regression equation between radium concentration and uranium concentration has been found from the regression plot:

$$Y(U_3O_8) = 1.384 X (Ra(eU_3O_8)) + 5.862 \text{ with } R^2 = 0.976 \quad (2)$$

The linear regression plot of studied samples has indicated a correlation coefficient of 0.976 (Figure 2). This plot indicates the association of daughter product with significant enrichment of parent

S. No.	BH No.	No of sample (n)	U ₃ O ₈ (ppm)			Ra(eU ₃ O ₈) ppm			Av. ThO ₂ (ppm)	DF	No. of samples	
			Min	Max	Av.	Min.	Max.	Av.			Above u/c	below u/c
1	KPU/23	8	107	2875	774	73	2028	624	21	1.17	8	0
2	KPU/57	9	91	1344	293	57	536	158	19	1.64	0	9
3	KPU/68	5	139	420	255	84	372	181	26	1.47	3	2
4	KPU/76	7	210	3600	1333	132	2270	830	57	1.47	7	0
5	KPU/79	33	122	13529	1749	91	8185	1236	26	1.32	23	10
6	KPU/82	5	140	521	285	104	413	221	17	1.34	3	2
7	KPU/93	11	99	715	330	83	592	267	5	1.27	11	0
8	KPU/108	5	94	448	203	69	294	139	21	1.44	5	0
9	KPU/110	14	113	1848	488	111	996	303	32	1.37	0	14
10	KPU/112	28	95	2730	516	64	2483	375	22	1.35	28	0
11	KPU/118	31	95	1099	291	75	927	251	20	1.28	22	9
12	KPU/123	17	90	1247	352	59	941	265	34	1.4	17	0
13	KPU/124	40	92	7040	656	55	6034	530	22	1.29	16	24
14	KPU/131	27	90	1095	207	68	820	163	26	1.28	23	4
15	KPU/138	113	93	8933	1489	80	7689	1215	25	1.26	93	20
16	KPU/144	50	105	7298	724	78	5081	533	63	1.33	12	38
17	KPU/147	16	101	9851	1179	96	6881	795	78	1.33	3	13
18	KPU/152	31	111	2104	422	98	1541	335	12	1.23	21	10
19	KPU/153	13	123	4264	621	97	3104	470	6	1.27	13	0
20	KPU/155	6	117	3055	1154	85	1908	781	12	1.37	6	0
21	KPU/156	16	122	559	226	94	302	160	11	1.44	16	0
22	KPU/163	65	91	7863	841	72	4635	594	27	1.33	28	37
23	KPU/166	42	95	6700	558	61	5098	402	27	1.44	22	20
24	KPU/169	12	229	3165	1048	187	2188	712	18	1.36	12	0
25	KPU/172	26	102	3241	802	81	2297	566	9	1.4	20	6
26	KPU/180	28	110	13851	2004	28	6885	1333	15	1.88	28	0
27	KPU/181	17	105	4973	1506	15	3491	1170	33	1.62	16	1
28	KPU/184	66	114	6066	509	63	2240	307	15	1.5	52	14
29	KPU/226	7	95	681	316	100	373	224	13	1.32	7	0
30	KPU/230	10	118	3623	632	84	1968	507	8	1.43	10	0
31	KPU/242	13	101	9558	3071	86	5652	2065	6	1.29	13	0
32	KPU/247	55	115	3923	544	82	2909	394	16	1.59	45	10
33	KPU/252	16	90	953	259	46	358	129	32	2	9	7
34	KPU/255	28	90	907	358	74	679	245	28	1.46	28	0
	Total	870										
	Average		112	4120	765	80	2711	544	24	1.41	620	250

Table 2: Details of U₃O₈, Ra(eU₃O₈) and disequilibrium factor (DF) of borehole core samples of Koppunuru area, Guntur district, Andhra Pradesh.

and good correlation among them. The average ThO₂ concentrations of these borehole core samples range from 5 ppm to 78 ppm. The split details of number of boreholes, the concentration ratio of ThO₂/ U₃O₈ in different boreholes is shown in Figure 3. This histogram is showing statistics of the distribution pattern of boreholes in different ratio ranges of ThO₂/ U₃O₈ viz. maximum 10 boreholes fall between 0.01-0.02 range while only 1 borehole falls under 0.13-0.14 range category.

The disequilibrium factor (DF) in the sample is calculated by the following formula

$$DF = \frac{U_3O_8 \text{ in the sample}}{Ra(eU_3O_8) \text{ in the sample}} \quad (3)$$

The disequilibrium is towards the parent uranium if the value of DF is more than one (DF>1). This is favourable for the prospector as it shows enrichment of uranium resulting in positive corrections in the final ore reserve estimation based on total gamma ray logging data. In contrast, if the value of DF is less than one (DF<1), then disequilibrium

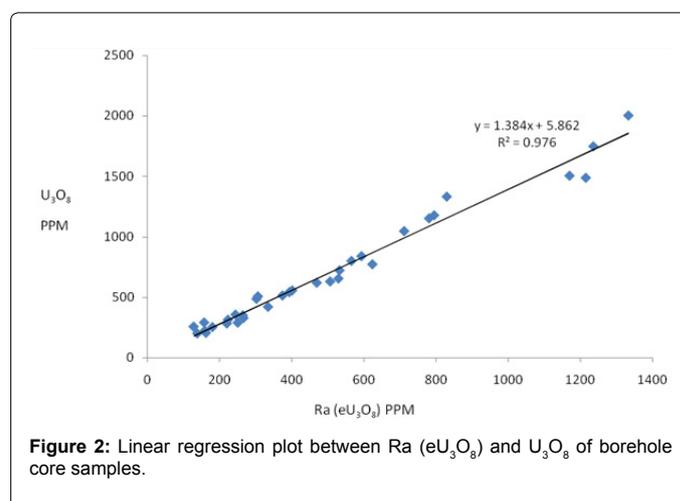


Figure 2: Linear regression plot between Ra (eU₃O₈) and U₃O₈ of borehole core samples.

is towards the daughter radium and is a non-desirable condition for uranium prospecting. It signifies partial removal of uranium from the system leading to a lowering of final ore reserve estimates based on total gamma ray logging data. All the boreholes have shown average DF significantly greater than 1 and the disequilibrium factors for studied core samples are listed in Table 2, which shows an average value of 1.41. This suggests enrichment of uranium either due to remobilization and deposition of uranium at the present locale or leaching of daughter products of the uranium series leading to an increase in concentration of parent uranium. These features are further supported by the presence of fractures, faults, felsic and mafic intrusive signifying pre- and post-depositional reactivation in the area providing a hydrothermal gradient for remobilization [17,19]. In addition, the presence of higher hydrouanium content (<10 ppb) away from the ore deposit suggests a possible role of groundwater on radioelement migration and fixation at suitable locales [27].

The studied mineralised core samples are broadly classified in two groups i.e. granite (below uniformity) and cover rock of Banganapalle quartzite/grit (above the unconformity). The disequilibrium factor is separately calculated for both types of samples. Average disequilibrium factor for the granite and quartzite/grit samples is 1.41 (Table 3) and 1.40 (Table 4) respectively. Thus, the study distinctly indicates close similarity in disequilibrium factor for both groups of samples, irrespective of the different lithic compositions and geologic position.

Impact on the ore reserve estimation

The presence of disequilibrium in uranium series between parent and daughter Radium-226 implies that ore grades of mineralised zones demarcated based on total gamma ray logging results needs to be corrected. Usage of Disequilibrium factor will lead to an upward revision in the total estimated resources. The disequilibrium correction factor calculated for core samples of different boreholes ranges from 1.29 to 1.47 of eU_3O_8 (Table 5).

Conclusion

The disequilibrium studies on the mineralised samples from the boreholes of the Koppunuru deposit has indicated:

1. Presence of strong disequilibrium in favour of parent uranium, with average value of 1.41.
2. DF is nearly same in the mineralisation hosted by both basement granites and Banganapalle quartzite.
3. Presence of disequilibrium in the mineralised zone implies an

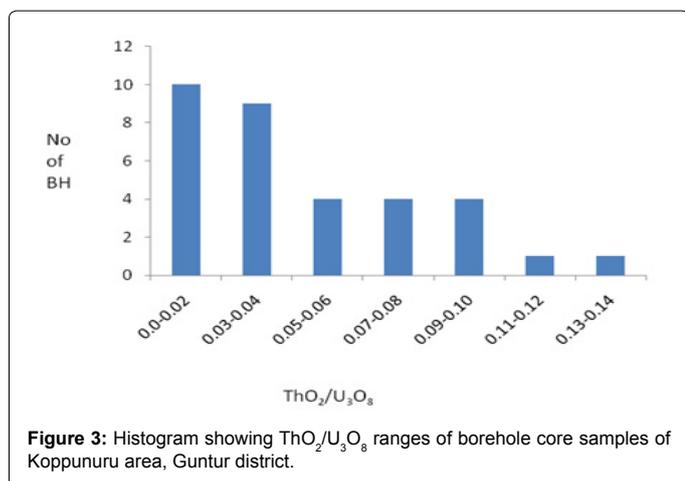


Figure 3: Histogram showing ThO₂/U₃O₈ ranges of borehole core samples of Koppunuru area, Guntur district.

S. No.	BH. No.	No. of samples	DF
1	KPU/57	9	1.64
2	KPU/68	2	1.06
3	KPU/79	10	1.2
4	KPU/82	2	1.57
5	KPU/110	14	1.37
6	KPU/118	9	1.36
7	KPU/124	24	1.29
8	KPU/131	4	1.28
9	KPU/138	20	1.46
10	KPU/144	38	1.32
11	KPU/147	13	1.33
12	KPU/152	10	1.36
13	KPU/163	37	1.34
14	KPU/166	20	1.49
15	KPU/172	6	1.62
16	KPU/181	1	1.03
17	KPU/184	14	1.45
18	KPU/247	10	1.86
19	KPU/252	7	1.85
		Total=250	Average=1.41

Table 3: Disequilibrium factor of Granites borehole core samples (below unconformity), Koppunuru area, Guntur District, Andhra Pradesh.

S. No.	BH. No.	No. of samples	DF
1	KPU/23	8	1.17
2	KPU/68	3	1.75
3	KPU/76	7	1.47
4	KPU/79	23	1.38
5	KPU/82	3	1.18
6	KPU/93	11	1.27
7	KPU/108	5	1.44
8	KPU/112	28	1.35
9	KPU/118	22	1.25
10	KPU/123	17	1.4
11	KPU/124	16	1.3
12	KPU/131	23	1.28
13	KPU/138	93	1.24
14	KPU/144	12	1.36
15	KPU/147	3	1.36
16	KPU/152	21	1.2
17	KPU/153	13	1.27
18	KPU/255	6	1.37
19	KPU/256	16	1.44
20	KPU/163	28	1.3
21	KPU/166	22	1.39
22	KPU/169	12	1.36
23	KPU/172	20	1.33
24	KPU/180	28	1.88
25	KPU/181	16	1.66
26	KPU/184	52	1.51
27	KPU/226	7	1.32
28	KPU/230	10	1.43
29	KPU/242	13	1.29
30	KPU/247	45	1.51
31	KPU/252	9	1.79
32	KPU/255	28	1.46
		Total=620	Average=1.40

Table 4: Disequilibrium factor of Banganapalle quartzite /grit borehole core samples (above unconformity), Koppunuru area, Guntur district, Andhra Pradesh.

Range eU ₃ O ₈ (ppm)	No of Samples	Avg. U ₃ O ₈ (ppm)	DF=U/Ra
90-200	412	162	1.47
201-400	204	332	1.29
401-600	49	613	1.32
601-800	33	953	1.39
801-1000	20	1264	1.45
1001-1500	53	1711	1.4
1501-2000	28	2275	1.33
2001-2500	21	3067	1.42
2501-3000	9	3480	1.3
3001-4000	18	4572	1.35
4001-8500	22	7817	1.39

Table 5: Disequilibrium factor core samples with different eU₃O₈ ranges of Koppunuru area, Guntur district, Andhra Pradesh.

upward correction in the ore grades, calculated based on total gamma log, by a factor of 1.41. Thus, an increase in grade and tonnage of the total estimated resources in Koppunuru deposit is indicated.

Acknowledgment

The authors are grateful to Shri. Parihar PS, Director, AMD, Hyderabad for giving permission to publish this paper. They also extend sincere thanks to Shri. Nanda LK, Additional Director (OP-III) and Dr. A.K. Chaturvedi, Additional Director (R&D), AMD, Hyderabad for their suggestions and encouragement.

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