

Paleocurrent Analysis of the Sandstone Body in Akpoha and its Enviorns (Lower Benue Trough) South-East Nigeria

Akana TS* and Didei IS

Department of Geology, Niger Delta University, Amassoma, Bayelsa State, Nigeria

Abstract

Palocurrent analysis study was carried out in Amsiri, Akpoha and its environing in the lower Benue Trough. This study was carried out to determine the direction of ancient flow. Compass/clinometers were used to obtain several measurements from outcrops within the study area. Sedimentary structures were also observed in the study area which includes tabular cross stratification, beddings, herribone structures, asymmetrical ripples and trough cross beds. Asymmetrical ripples and trough cross beds indicates a low energy current. Tabular cross stratifications, herribone structures, where formed by high energy currents. The herribone structures are evidence of different paleocurrent directions as they are formed when current periodically flows in opposite directions. Based on the Rose diagram plot, the paleocurrent direction pointed at two directions; NE direction in location 2 and 3 which are Ameata and Ibii and the NE/SW/NW direction in location 1 which is Akpoha. This however shows that the sediments were derived from two main sources.

Keywords: Sandstone; Amasiri; Paleocurrent; Sedimentary structures; Direction

Introduction

Paleocurrent study is an important part of sedimentology; it provides direct information about the orientation of the sedimentary system. Paleocurrent analysis can yield information about the flow direction of rivers, long shore currents, sediments, gravity flow and paleowinds. This in turn helps establish a model of paleogeography, which can lead to better predictability of economically important facies. The work studies directional and other primary properties of sedimentary rocks from which we can reconstruct current systems and make a paleocurrent analysis.

The examination of modern seas, estuaries, and rivers, shows that there is a distinct relationship between their physical and the current present in them; currents so impress themselves on the deposit formed under their influence that their characters can be ascertained from those formed in the ancient periods. Therefore their physical geography can be inferred within certain limits. Paleocurrent analysis has had a long heritage that includes a large portion of geology. It is closely related to and dependent on stratigraphy, sedimentary petrology and structural petrology.

Paleocurrent analysis can aid in the prediction of both regional and local sedimentary trends in exploration for petroleum and ground water, can predict the trend of "washouts" or "cut-outs" in coal fields and can aid in mineral exploration where the deposit are related to sedimentary trends as are some gold and uranium deposits. In stratigraghy, paleocurrent analysis can contribute to improved regional correlation by establishing depositional strike. Paleocurrent analysis depends on evidence obtained from the outcrop, the sub surface and the laboratory. Some studies may rely only on field methods, some may use only laboratory evidence, and others combination of all these three for integrated paleocurrent analysis of most sedimentary basins. In any case, however, a distinguishing feature of paleocurrent analysis places its emphasis upon quantitative study. This aspect of paleocurrent analysis is a further development of the accelerating quantification of geologic data.

Paleocurrent analysis depends on the presence in sedimentary

rocks of numerous properties from which current direction can be inferred. These properties occur, in varying kinds and abundant, in all sedimentary basins from the present to the far geologic past.

The Benue Trough is an intercontinental rift basin characterised by tectonic and magnetic activities during cretaceous time. The Benue Trough can be sub-divided into lower, middle, and upper parts that were affected by santonian compression which deformed the Benue Trough by inverting the main depocenter of the Abakaliki Trough and subsequently creating the Anambra and Afikpo basin to the north-west and south-east respectively [1,2]. The santoniantectonism differentiated the sedimentary succession into pre and post santonian packages. The post santonianmaastrichtian deposit occur both in the Anambra and Afikpo basin respectively [3]. The Nkporo formation rest on the post santonian unconformity plane which lies above the pre santonianEze-Aku formation of Turonian times. The Asu River Group rest on the basement and lies below the Eze-Aku formation.

Location and accessibility of study area

The study area covers an area within longitude 7° 50°E and 7° 55°E, and latitude 5°50°N and 5°55°N. The area is also located in the Afikpo syncline in South Eastern Nigeria which is in Afikpo local Government area of Ebonyi Nigeria. It covers an area of about 60 km² (sixty square kilometres) (Figure 1).

Geographic coordinates of Afikpo, Nigeria

Latitude: 5°50 State '35" N Longitude: 7°50'14", Elevation above sea level: 83 m = 272 ft

*Corresponding author: Akana TS, Department of Geology, Niger Delta University, Amassoma, Bayelsa State, Nigeria, Tel: +234809757544; E-mail: innosilicate@yahoo.com

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The vegetation of the Afikpo area belongs to the Guinea Savannah vegetation belt of Nigeria which is associated with thick vegetation cover similar to that of Rain Forest, as part of the south east Nigerian tropical climate condition (Figure 2). This vegetation cover is more pronounced in the flood plain area where they are more nutrient rich in soil is characterized by tall trees, climbers, shrubs with abundant palm trees and some areas where there are no or less nutrient, where by the trees find it difficult to survive. The vegetation is influenced by human

In Akpoha, in Afikpo there are two main types of soil present, the clayey hydromorphic soil and the grey sandy clay hydromorphic soil. There are two types of superficial deposit, they are Laterites and Alluvium (Figure 4).

activities, annual rainfall, topography and other climatic factors (Figure 3).







Figure 3: The figure shows the vegetation and topography of Akpoha and its environs.



Regional Geology

The evolution of sedimentary basins in South-Eastern Nigeria followed the opening of the South Atlantic and the break-up of the South American and West African plates in late Jurassic times. The probasin of the Benue Trough was the failed arm of an RRR triple junction which extended from the Northern limits of the Niger Delta Basin to the Chad area in the NE [4].

A review of the tectonic framework of the inland sedimentary basin of Nigeria has demonstrated that the Benue Trough system was indeed a rejuvenation of existing basement fractures. Wrench movements along these faults resulted in blocks faulting and formation of several sedimentary basins. The Benue Trough system had in the past been conveniently subdivided into the lower, middle and upper Benue Trough. However recent aeromagnetic and gravity data across the entire system has demonstrated the distinct nature of these basins each with its well defined sedimentary succession and separated by positive anomaly areas. Sediment thicknesses of up to 8 km have been recorded in some of these basins within the trough. In South-eastern Nigeria they include the Abakaliki, Anambra and Afikpo basins, as well as other minor basins such as the Calabar Flank and Mamfe Embayment which together form the lower Benue Trough. The Paleogene basin has been influenced mainly by the later generation transform faults (Charcot,

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Chain, Ascension & St. Paul's) which emanate from the mid-Atlantic ridge system and now extend only into the Tertiary basin. Up to 12 km sediments have been deposited within the central part of the Niger Delta basin, much of which is above oceanic basement (Figure 5).

The Benue Trough has often been described as elongated, partly fault bounded depression occupied by up to 6000 meters marine, fluvial and deltaic sediments that have been compressionally folded in a nonorogenic shield environment [5] emphasized the wholly nature of the Trough, the thin crust beneath it and the dominate influence of vertical movement as a course of the folding.

Tectonic setting

An adequate interpretation of the geology and history of the Benue Trough requires a proper understanding of its regional tectonic framework. The origin, tectonic evolution, and the stratigraphy of the Benue Trough has been discussed by several authors [1-7] considered the Benue Trough as a continuation of a sinistral shear fracture zone crossing the Mid Atlantic Ridge with a counterpart in Brazil. Wright [6], considered the trough as a sediment filled rift of tensional origin. Burke [4] and Grant [7] on their part, considered the Benue Trough as a "failed" branch of a triple plate margin junction. Burke et al, 1970 went further to propose an aulacogen model which refers to the Benue Trough as the failed arm of a triple RRR (rift-rift) junction which evolved by mantle upwelling, crustal stretching and the separation of the West African South American continental plates. The gravity profile across the Benue Trough, which is remarkably similar to that of the Red Sea and the presence of igneous intrusion of alkaline genetic class seem to support the aulacogen model. Though this model seems to be very popular, it failed to account for the Santonian compressive phase in the evolution of the Benue Trough (Figure 6) [4].

Three mega sequences: Asu River Group, Eze-Aku Group and the proto-Niger Delta sediments have been intruded by basic rocks during the three tectonic events, namely Cenomanian, Santonian and Maastrichtain times [8,9].

Geology of the study area

In the Afikpo Basin, three main Cretaceous lithostratigraphic units have been recognized, namely; The Asu River Group, the Eze-Aku Group and the post Santonian Proto-Niger Delta succession [8]. The area consists of some parallel and high anticline and syncline structures that show outcrops.

The basal sediments are 3000 m thick Asu- River Group (Late Albian – Early Cenomanian) consisting of the Awi Formation, Abakalikishales



South-Eastern sedimentary basins (Odigi, 2007).

STAGES & EPOCHS		LOWER BENUE	MIDDLE BENUE		UPPER BENUE		LOWER BENUE TROUGH
		ANAMBRA BASIN	LAFIA AREA	BASHAR AREA	GOMBE AREA	LAU AREA	AFIKPO BASIN
(BRIDE)	Eocene	Ameki Fm.					Ameki Fm.
	Paleocene	Imo Shale	Volcanics	Kerri Kerri Fm	Kerri Kerri Fm	Volcanics	Imo Fm
		Nsukka Fm.		~~~~~~~~~~~	1	[Nsukka Fm.
MAESTRICHTIAN		Ajali Sandstone		Gombe Sandstone	Gombe Sandstone	Lamja Sandstone	Ajali Fm.
		Mamu Formation	Lafia Formation			Shale	Mamu Fm.
HOMMAN	Campanian	Enugu Shale		~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~		Numarina	Nkporo Fm
	Santonain	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~		Unnamed Marine	Pindiga Formation	Sekule Formation	
	Coniacian	Awgu Formation	Awgu Formation			Dukul Formation	1
the	Upper	For Alm Ohele		TEomation	Yolde Formation	Yolde Formation	Eze Aku Group
Rom	Lower	Eze Aku Shale	Eze Aku Fm.	Zurak rome	Sandstone	Sandstone	E20 AND GIVEP
CENOMANIAN		Odukpani Fm.	Keana Formation	Keana Fm	Bima our	Bima Sanoo	
ALBIAN Upper		Asu River Group	Asu Awe Formation	Pre-Bima Sediment	Pre-Bima Sediment	Pre-Bima Sediment	Asu River Group
PRECAMBRIAN		Basement Complex					Basement
		HOQUE (1977)	OFFODILE (1976)	AYOOLA (1978)	CARTER ET AL (196 CRATCHELY & JON	3), ES (1965)	ODIGI, 2007

and the Awe Formation. Unconformably overlying the Asu River Group are 200 m thick Late Cenomanian – Early Santonian sediments of the Eze-Aku Group. The Eze-Aku Group is unconformably overlain by post-Santonian sediments of the pro-Niger Delta beds [9-11].

The oldest sedimentary rocks overlying the pre-Cambrian Oban massif basement complex rocks are non-marine to marginal marine deposits. The Asu River Group comprises of sandstones and dark shales with Ammonites. The Asu River Group represents deposits of the first transgressive-regressive marine depositional cycle in the area [12].

The Eze-Aku beds overly the Asu river Group unconformably and consists of shales, limestones and sandstones ridges which strike, on the average, NW 40° SE with dip ranging from 20° to 68° [13,14]. The sandstone bodies are parallel to elongate features. The Eze-Aku beds represents the second transgressive depositional cycle that occured during the Upper Cretaceous [1].

The end of the Turonnian was marked by a regressive phase which was terminated by a third sedimentation cycle of the Campanian-Maestrichian transgression and regression which led to the deposition of the post-SantonianCampanian-Maestrichian proto-Niger Delta Sequence in the broad and gentle depression of the Afikpo Basin. The post-Santonian (Campanian-Maestrichian) sediments overly the Eze-Aku strata unconformably [15]. These beds comprises of alluvial, fluvialshallow marine conglomerates and sandstone, mudstone and shale which show a NE - SW trend. Following the Southwestward Shift of the depositional axis in the Southern Benue trough, the parallicNkporo Formation, Mamu, Ajali and Nsukka Formations were deposited in the subsiding Afikpo basin [16]. The Late CampanianNkporo Formation consists of basal alluvial sediments and conglomerates overlain by alteration of sandstones, siltstones, shales, mudstones, and coal beds (0.250 - 5 m thick). Basin ward, these sequences laterally changes to deltaic and tidal channel deposits.

Methods and Materials

This has to do with various methods employed in the study. The method applied in the collection of data in the course of carrying out this research work involved several methods. The methodology applied in this study encompasses description of outcrop sections of the Amasiri sandstone exposed behind Government College in Afikpo area.

The method employed in this research work was field studies. This was done in order to obtain data (results) on paleocurrent and sedimentary structures occurring Amasiri sandstone; behind

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Government Secondary School Afikpo and Government technical College Akpoha [17-20].

At the exposure, Sedimentary structures, textures and Colour, orientation patterns of sediments were observed. Field Studies include taking photographs of sedimentary structures and also measurement of strike and dip were also taken with a compass clinometer which would help in showing paleocurrent direction [21].

The field method was used specifically for the collection and interpretation of data behind government college, Afikpo.

Measurements were done at regular interval (10 ft) from the base of the outcrop to the top in a vertical section [22-24]. The geological field mapping of this work involved description, and sampling of the exposed section of the sandstone of Eze-Aku formation. The sandstone outcropping behind Government College, Afikpo were studied.

Results and Discussion

Data presentation

This data was obtained from the field using a compass clinometer measuring from outcrop to outcrop in different locations and paleocurrent analysis measurements are acquired from deformed beds.

Paleocurrent analysis (Rose diagram)

Location 1: 282°, 286°, 252°, 248°, 228°, 256°, 238°, 256°, 212°, 262°, 280°, 300°, 258°, 252°, 308°, 324°, 238°, 244°, 274°, 238°, 232°, 236°, 270°, 286°, 320°, 328°, 288°, 316°, 306°, 292°, 40°, 44°, 22°, 30°, 20°, 24°, 58°, 46°,66°, 42°, 88°, 42°, 22°, 26°, 78°, 24°, 50°, 38°, 40°, 66°, 44° (Figure 7) (Table 1).

Location 2: 84, 62, 68, 88, 76, 60, 62, 64, 78, 18, 54, 48, 72, 76, 60, 20, 84, 80, 82, 68, 64, 62, 80 (Figure 8) (Table 2).

Location 3: 60°, 50°, 42°, 40°, 66°, 36°, 84°, 34°, 70°, 72°, 88°, 75°, 63°, 80°, 46°, 55°, 67°, 77°, 80°, 84°, 67°, 50°, 46°, 53°, 70°, 32°, 56°, 37°, 40°, 80°, 75°, 77°, 60°, 53°, 72°, 39°, 49°, 59°, 69°, 79°, 89° (Figure 9) (Table 3).

Stereonet

From the Rose diagram plotted, the current direction is interpreted to be in the NE direction in locations two and three which are Ameata and and Ibii town. This implies that the formations in this area were deposited in NE direction. While for that of location one (1) Akpoha



Class Interval	Tally	Frequency
20-90	++++ ++++ ++++	21
90-250	++++	9
250-300	++++ ++++	14
300-360	++++-	6

Table 1: Frequency table for plot 1 location 2.



Class Interval	Tally	Frequency
10-50		3
50-90		20

 Table 2: Frequency table for plot 2.



Technical College, the paleocurrent direction as measured from the inclination of the cross beds is pointing to the NE/SW/NW direction. From the differential directions of the paleocurrent, it can be said to

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have resulted from a reversal of current, as the current that deposited sediments in the area are bimodal (predominant current directions) and bipolar (two opposed direction) (Figure 10).

The paleocurrent direction is NE in Ameata and Ibii Town. The paleoflow is North-eastern, this suggesting that the sediments were derived from the Oban Massif, in the Calabar flank which is located in the Northeast. The Amasiri Sandstone together with the adjoining Konshisha River Group deposits were largely derived from the Basement Complex areas around Ogoja [25].

The paleocurrent in the NE/SW/NW direction shows that sediments depositedcame from the Oban Massif which is NE, the Anambra basin sediments came from the NW direction, and the SE

Class Interval	Tally	Frequency	
30-50	++++ ++++	11	
50-70	++++ ++++	14	
70-90	++++ ++++ ++++	15	

Table 3: Frequency table for plot 3.



and SW deposition indicates that sediment where deposited from transgression and regression of the sea phase.

The SW directed paleoflow is measured from the friable, cross bedded, arenaceous sandstones. Since the paleoflow is in southwesterly direction, an indication shows that it is from Basement Complex areas around Ogoja.

Depositional Environment

The grain size of sandstone outcrops ranges from fined to medium grained quartz arenite, they are poorly or moderately sorted. Based on this parameter, the formation indicates low energy and moderate energy depositional medium.

A combination of sedimentary structures is strong indicators of a particular sedimentary environment and at the study area; beddings, herringbone structures tabular cross beds etc. are common. These sedimentary structures are as a result of low to high energy current.

Sedimentary Structures

Based on the types of sedimentary structures found in the study area, paleocurrent of the area can be inferred. Sedimentary structures observed includes beddings, tabular cross stratification, cross lamination, herringbone structures, asymmetrical ripples, trough cross beds

Asymmetrical ripples and the trough cross beds signify the lower part of the lower flow regime i.e. they were formed by low energy current. Tabular cross stratification, herringbone structures were formed by high energy current that deposited the sediment. The herringbone structures are evidence of differential paleocurrent directions, as they are formed when current periodically flows in opposing direction. Bedding planes of the formation shows time-medium strata, which is as a result of alternating high and low energy deposition. A high energy current tends to deposit coarser sediments than a low energy current.

Conclusion

This work has examined the sedimentology of the Amasiri sandstone of the Eze-Aku Group exposed at the eastern flank of Abakaliki fold belt with respect to the fluid of observation gathered and also gotten from previous works on the geology of the southern Benue Trough. The Benue Trough was formed as a result of series of tectonism and repetitive sedimentation. A clean imprint of this information had been left in rocks found in the area. Highly undulating alteration of sandstone ridges and shale low land trending in a NE-SW direction are found in the area, and also some sedimentary structures were found on the surface of the rock. The Amasiri sandstone is moderately sorted, very fine to coarse grained, highly consolidated and feldspathicarenite sandstone. Marine fossils can be found in them. A combination of sedimentary structures is strong indicators of a particular sedimentary environment and at the study area. The sedimentary structures are as a result of low to high energy current. From the Rose diagram the currents direction is interpreted to be in the NE direction and NE/SW/ NW direction at three different locations. This implies that there was a reversal of current, as current that deposited sediments are bimodal (predominant current directions) and bipolar (two opposed direction). The paleoflow is North-eastern, this suggesting that the sediments were derived from the Oban Massif, in the Calabar flank which is located in the Northeast. The Amasiri Sandstone together with the adjoining Konshisha River Group deposits were largely derived from the Basement Complex areas around Ogoja. The sediments might have

been derived from the North eastern part of Nigeria (younger granites), south eastern basement complex and the Abakaliki syncline.

References

- 1. Murat RC (1972) Stratigraphy of Paleogeography of the Cretaceous and Lower Tertiary in Southern Nigeria. African Geology.
- Benkhelil J (1989) The origin and Evolution of Cretaceous Benue Trough, Nigeria. African Earth Sciences J 8: 251- 282.
- 3. Reyment RA (1965) Aspect of the Geology of Nigeria. University Press, Ibadan, Nigeria.
- Burke KC (1970) Geological history of the Benue valley and adjacent areas. (Ed) African Geology university of Ibadan press, Ibadan pp. 187 – 202.
- Olade MA (1975) Evolution of Nigeria Benue Trough (Aulocogen): A tectonic model. Geol Mag 112: 576- 583.
- Wright JB (1989) Review of the Origin and Evolution of the Benue Trough in Nigeria. In: Kogbe C. A. (ed). Geology of Nigeria, Elizabeth publication company, Lagos, 359-364.
- Grant NK (1971) TheSouth Atlantic, Benue trough and Gulf of Guinea cretaceous triple junction. Geological Society of America Bulletin 82: 2295 – 2298.
- 8. Odigi MI (2007) Stratigraphic Succession in the lower Benue Trough.
- Odigi MI, Amajor LC (2008) Petrology and Geochemistry of Sandstones in the Southern Benue Trough of Nigeria. Implication to Provenance and Tectonic Setting. Chinese Journal of Geochemistry. 27: 384-394.
- 10. Allen JRL (1965) On Bedforms and Paleocurrents Sedimentology. 6: 153-189.
- 11. Allen JRL (1982) Sedimentology structues, their characteristics and physical basic, Development Sedimentology 30 A and B, 1: 539, 11: 663, Elsevier science, New York.

- 12. Collinson JD, Thompson DB (1982) Sedimentary structure 2nd Edition. Chapman and Hall, London, 207 pp.
- Dalrymple RW (1992) Tidal depositional system. In: RG Walker and James NP (Eds) Facies Models: Geol Association Canada pp. 201.
- 14. Etu-Efeotor JO (1998) Stratigraphy, sedimentology and depositional environment of reservoirs sands of the Ivol Field Niger Delta. Glob Jour Pure and Applied science Vol 4.
- Inyang PEB (1978) The Climatic Regions. In: Ofomata G. E. K. (Ed). The Nsukka Environment. Ethiope Publication, Benin. Pp.17 – 24.
- Kogbe CA (1976) Cretaceous and paleocene sediments of Southern Nigeria. In: Kogbe CA (Ed) Geology of Nigeria. Elizabeth publ co Lagos pp. 273-282.
- 17. Kogbe CA (1975) Paleogeography History of Nigeria from Albiam.
- Kogbe CA (1989) Geology of Nigeria 2nd Revised Edition Rock View International, Tour Onyx 10 Rue Vandrezanna.
- 19. Krumbein WC, Sloss L (1963) Stratigraphy and sedimentation. Pp.234.
- 20. Miall AD (2000) Principle of Sedimentary Basin Analysis. Springer Verlag, Berlin Heidelberg, pp.616.
- Maurice E, Tuker F (1985) Sedimentary Petrology and Introduction of Origin of Sedimentary Rocks.
- Nwachukwu SO (1972) The Tectonic Evolution of the Southern Portion of the Benue Trough, Nigeria. Geol Mag 109: 411 – 419.
- Ogbonnaya CO (2012) Sedimentary and Facies Analysis of Amasiri sandstone, Lower Benue Trough. South Eastern Nigeria. Pp.6 & 138.
- 24. Reyment RA, Barber WM (1965) Nigeria and Cameroon. In: Lexique Stratigraphic International: Afrique 4: 35-39.
- Tattam CM (1944) A Review of Nigeria Stratigraphy. Research and Educational Development of the Geological Survey of Nigeria pp.27-46.

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